

For Reference

NOT TO BE TAKEN FROM THIS ROOM

Ex LIBRIS
UNIVERSITATIS
ALBERTAENSIS



THE UNIVERSITY OF ALBERTA

RELEASE FORM

NAME OF AUTHOR ZHENG BAO-SHAN
TITLE OF THESIS SOME PROBLEMS INVOLVING THE RAY
 APPROXIMATION OF THE WAVE EQUATION
DEGREE FOR WHICH THESIS WAS PRESENTED MASTER OF SCIENCE
YEAR THIS DEGREE GRANTED SPRING 1985

Permission is hereby granted to THE UNIVERSITY OF ALBERTA LIBRARY to reproduce single copies of this thesis and to lend or sell such copies for private, scholarly or scientific research purposes only.

The author reserves other publication rights, and neither the thesis nor extensive extracts from it may be printed or otherwise reproduced without the author's written permission.

THE UNIVERSITY OF ALBERTA

SOME PROBLEMS INVOLVING THE RAY APPROXIMATION
OF THE WAVE EQUATION

BY

©
ZHENG BAO-SHAN

A THESIS

SUBMITTED TO THE FACULTY OF GRADUATE STUDIES AND RESEACH
IN PARTIAL FULFILMENT OF THE REQUIREMENTS FOR THE DEGREE
OF MASTER OF SCIENCE

IN

GEOPHYSICS

DEPARTMENT OF PHYSICS

EDMONTON, ALBERTA

SPRING 1985



Digitized by the Internet Archive
in 2026 with funding from
University of Alberta Library

<https://archive.org/details/0162006822207>

THE UNIVERSITY OF ALBERTA
FACULTY OF GRADUATE STUDIES AND RESEARCH

The undersigned certify that they have read, and recommend to the Faculty of Graduate Studies and Research, for acceptance, a thesis entitled SOME PROBLEMS INVOLVING THE RAY APPROXIMATION OF THE WAVE EQUATION submitted by ZHENG BAO-SHAN in partial fulfilment of the requirements for the degree of MASTER OF SCIENCE in GEOPHYSICS.

Abstract

In this paper the conditions under which the equation of motion in a 3-D inhomogeneous isotropic medium can be approximated by the scalar wave equation, and the solution to the scalar wave equation approximated by rays are investigated. The relationship between these conditions is also discussed.

Accurately and efficiently solving the eikonal equation which describes the kinematic characteristics of the wave motion is one of the more important problems in ray approximation methods. It is convenient to solve the eikonal equation as a boundary value problem in some geophysical problems. Much attention is given to the investigation of Pereyra's method of solving a boundary value problem, which is a variable order finite difference method on nonuniform meshes. These meshes are selected automatically by the program as the computation proceeds. We have programed this method for 2 and 3 dimensional velocity models with linear gradients. This program is tested for models with linear velocity gradients in one and two space variables. This method yields satisfactory results when compared with known analytic and numerical solutions given by W. H. K. Lee's programs.

A new method based on the numerical solution of the eikonal equation is presented in this thesis.

ACKNOWLEDGEMENTS

I wish to express my many thanks to Dr. F. Hron for the suggestion of these topics and for his assistance in the completion of the study.

I am grateful to Dr. P. F. Daley and Mr. Roger Chan for helpful discussions.

During the course of this study, I received financial support from the Department of Physics, the University of Alberta.

List of Tables

Table	Page
1. Solutions of the example about the nonuniform mesh ..	40
2. Comparison of analytic, SRT1 and SRT2 solutions for case No.1	44
3. Computed solution of ZH1 for case No.1	46
4. Comparison of SRT1 and SRT2 solutions for case No.2..	49
5. Computed solution of ZH1 for case No.2	51
6. Comparison of analytic, SRT1, SRT2 and ZH1 solutions	54
7. Main computed solutions using ZH1 program	56
8. Computed solution of ZH1 for model 1	57
9. Computed solution of ZH1 for model 2	58
10. Computed solution of ZH1 for model 3	59
11. Computed solution of ZH1 for model 4	60
12. Computed solution of ZH1 for model 5	61

Table of Contents

Chapter	Page
1. Introduction	1
2. The Wave Approximation to the Equation of Motion	4
3. The Ray Equation Approximation of the Wave Equation	8
4. The Energy (Amplitude) Equation	11
5. Discussion on the Relationship of the Approximating Conditions	17
6. The Two Point Seismic Ray Tracing Problem in an Inhomogeneous Medium	20
6.1 The Derivation of the Ray Path Equation	20
6.2 Two Point Seismic Ray Tracing Equation in 3-D Inhomogeneous Media	23
6.3 Two Point Seismic Ray Tracing Equation in 2-D Inhomogeneous Media	26
6.4 An Alternative Formulation of the Two Point Ray Tracing Problem in 2-D Media	28
6.5 Treatment of Interfaces in the Medium	29
6.6 A Simple Example for the Treatment of the Interfaces in a 2-D Medium	32
7. A Numerical Method for Solving Ray Equations	35
8. Computer Programs for Two Point Seismic Ray Tracing and Some Examples	41
8.1 Test Case No.1	41
8.2 Test Case No.2	48
8.3 Summary of Test Case Results	53
8.4 The Program ZH1 and 5 Computed Examples of Ray Tracing Problems in 2-D Media	55
8.5 A Simple Computed Example for 2-D Model With an Interface	62
9. Method for Computation of Ray Amplitudes	64

10. Conclusion78

REFERENCES79

APPENDIX I84

List of Figures

Figure	Page
1. Curvilinear coordinate system on the ray	12
2. Example of the nonuniform mesh technique	39
3. Geometry for case No.1	43
4. Definition of the taking off angle β and the azimuth angle α	65
5. The ray under consideration OA and the perturbed ray OB	68
6. Explanation of various symbols for the computation of the quantity of BB'	70
7. Definitions of the angles $\alpha(AB)$ and α_i and determination of the coordinate Z_i	72
8. Definition of the angle α_i'	74
9. Definition of the angle γ_i and determination of the coordinates X_i and Y_i	76

1. Introduction

Problems of wave propagation play an important role in various branches of geophysics. But so far, because of the complexity of the Earth's crust and the uppermost mantle, no analytical solutions of the equation of motion are known. There are two main approaches to the investigation of wave fields in such complicated media: approximate methods and methods based on direct numerical solution of the elastodynamic equation (such as the finite difference method).

Approximate methods can be used to study the propagation of elastic waves in quite complicated structures. Let us mention here at least Whitham's wave theory and various types of high-frequency approximations, such as the ray method.

Ray theory has been extensively used in seismology (see Hron, F. 1984). The ray method can be applied to more complicated media, in which other approximate methods cannot be used. The major problem in the application of the ray method is that it represents only a high frequency approximation and as such has a number of serious restrictions. Nevertheless, at the present time it is the only method that is able to give us an approximate answer to many problems of seismic body wave propagation in complicated geological models.

It is, therefore, very important to know the conditions under which the ray method approximation can be used to compute the wave field.

One of objectives of this paper is to discuss the restrictions for using the ray method. This problem will be discussed in sections 2 through 5.

In order to study the inhomogeneous structure of the Earth, seismologists have developed several techniques to trace seismic rays. So far there are two methods for tracing seismic rays between two given end points in three dimensional media. One of these methods involves "shooting" the ray from one point with a given starting direction and then changing this starting direction until the ray emerges at the desired target. It is called the "initial value method". Another method involves "bending" the initial path between the end points until it satisfies the principle of stationary time. This is called the "boundary value method". Some authors, for example, Yang and Lee (1976), Julian and Gubbins (1977), Pereyra, Lee and Keller (1980) have recently presented a comparison of these two methods, concluding that the boundary value method is computationally faster.

As Wesson (1971, p.741) pointed out, tracing seismic rays between two end points is required in seismological applications, such as earthquake location (e.g., Engdahl and Lee, 1976) and the determination of the three-dimensional velocity structure under a seismic array (e.g., Aki and Lee, 1976; Aki et al., 1977).

The second task of this thesis is to investigate the two point seismic ray tracing problem using the boundary value method in sections 6 through 8. We have programed Pereyra's method and have computed results for some models. High accuracy and efficiency are achieved by using variable-order finite difference methods on nonuniform meshes which are selected automatically by the program as the computation proceeds. The variable mesh technique allows the program to "adapt itself" to the particular problem at hand, thereby producing more detailed computations where they are needed, as in the case of highly curved rays.

The geometrical spreading coefficient which determines the amplitude of the ray may be retrieved easily from the "bending" solution. We will describe a new numerical method to calculate the amplitude of the ray in section 9.

2. The Wave Approximation to the Equation of Motion

The most general form of the equation of motion in a three dimensional inhomogeneous isotropic elastic medium may be written as follows:

$$(1) \quad \rho \partial^2 \vec{u} / \partial t^2 = (\lambda + \mu) \nabla \theta + \mu \nabla^2 \vec{u} + \theta \nabla \lambda + 2(\nabla \mu \cdot \nabla) \vec{u} + \nabla \mu \times (\nabla \times \vec{u}),$$

where λ and μ are Lames parameters, ρ is the volume density. These quantities are all functions of position. $\vec{u}$ is the displacement vector, and $\theta = \nabla \cdot \vec{u}$ is the volume expansion. Applying the rotation operator $(\nabla \times)$ and the divergence operator $(\nabla \cdot)$ in turn to (1) yields:

$$(2) \quad \rho \partial^2 \vec{\omega} / \partial t^2 + \nabla \rho \times \partial^2 \vec{u} / \partial t^2 = \mu \nabla^2 \vec{\omega} + \nabla \times [(\lambda + \mu) \nabla \theta] + \\ + \nabla \times (\theta \nabla \lambda) + \nabla \times (\nabla \mu \cdot \nabla \vec{u}) + \nabla \times (\nabla \vec{u} \cdot \nabla \mu),$$

$$(3) \quad \rho \partial^2 \theta / \partial t^2 + \nabla \rho \cdot \partial^2 \vec{u} / \partial t^2 = (\lambda + 2\mu) \nabla^2 \theta + \nabla \theta \cdot (\nabla \lambda + \nabla \mu) + \\ + \nabla \mu \cdot \nabla^2 \vec{u} + \nabla \lambda \cdot \nabla \theta + \theta \nabla^2 \lambda + \nabla \cdot (\nabla \mu \cdot \nabla \vec{u}) + \nabla \cdot (\nabla \vec{u} \cdot \nabla \mu),$$

where $\vec{\omega} = \nabla \times \vec{u}$ is the rotation of the displacement vector. In general, it is difficult to solve Eqs.(2) and (3) directly, although in many cases we are able to simplify them under some conditions and obtain an approximate solution of (1). Now we will consider these approximating conditions. We will study Eq.(3) using the similarity theorem of fluid dynamics (see Tritton, D. J. 1977). It is convenient to rewrite (3) as follows:

$$(4) \quad \rho \partial^2 \theta / \partial t^2 + \nabla \rho \cdot \partial^2 \vec{u} / \partial t^2 = (\lambda + 2\mu) \nabla^2 \theta + \nabla \theta \cdot \nabla (\lambda + \mu) +$$

$$+ \nabla \mu \cdot \nabla^2 \vec{u} + \nabla \lambda \cdot \nabla \theta + \theta \nabla^2 \lambda + (\nabla \nabla \mu) : \nabla \vec{u} + \\ + \nabla \mu \cdot \nabla (\nabla \cdot \vec{u}) + \nabla^2 \vec{u} \cdot \nabla \mu + \nabla \vec{u} : \nabla \mu \nabla,$$

where ":" indicates the dot product of dyadics. Let us introduce the characteristic dimensions for our problem as follows:

L characteristic length,

T characteristic time (period),

S characteristic quantity of displacement,

$\Theta = S/L$ characteristic quantity of volume
expansion θ ,

$U = L/T$ characteristic quantity of speed,

$$V = [(\lambda + 2\mu)/\rho]^{1/2}.$$

Then we have the following non-dimensional quantities of time t , position $\vec{X}$, speed V , volume expansion θ , and displacement vector $\vec{u}$:

$$(5) \quad t' = t/T,$$

$$(6) \quad \vec{X}' = \vec{X}/L,$$

$$(7) \quad V' = V/U,$$

$$(8) \quad \theta' = \theta/\Theta = \theta L/S,$$

$$(9) \quad \vec{u}' = \vec{u}/S.$$

Substituting Eq.(5)-(9) into Eq.(4) and using ∇' as the expression for the Laplacian relative to the coordinates $\vec{X}'$ and making some algebraic operations we have:

$$(10) \quad V'^{-2} \partial^2 \theta' / \partial t'^2 + V'^{-2} (\nabla' \rho / \rho) \cdot (\partial^2 \vec{u}' / \partial t'^2) = \nabla'^2 \theta' +$$

$$\begin{aligned}
& + (\lambda+2\mu)^{-1} \nabla'(\lambda+\mu) \cdot \nabla' \theta' + (\lambda+2\mu)^{-1} \nabla' \mu \cdot \nabla'^2 \bar{u}' + \\
& + (\lambda+2\mu)^{-1} \nabla' \lambda \cdot \nabla' \theta' + \theta' \nabla'^2 \lambda / (\lambda+2\mu) + (\lambda+2\mu)^{-1} \nabla' \nabla' \mu : \nabla' \bar{u}' + \\
& + (\lambda+2\mu)^{-1} \nabla' \mu \cdot \nabla' \theta' + \nabla'^2 \bar{u}' \cdot \nabla' \mu / (\lambda+2\mu) + \\
& + \nabla' \bar{u}' : \nabla' \mu \nabla' / (\lambda+2\mu).
\end{aligned}$$

Now let us discuss (10) instead of (3). Because $\nabla'^2 \lambda$ is of the same order as $\nabla' \lambda$; $\nabla' \nabla' \mu$ and $\nabla' \mu \nabla'$ as $\nabla' \mu$, by choosing a suitable S we can get $\partial^2 \theta' / \partial t'^2$ to have the same order as $\partial^2 \bar{u}' / \partial t'^2$, $\nabla' \theta'$, θ' , $\nabla'^2 \theta'$, $\nabla'^2 \bar{u}'$, $\nabla' \bar{u}'$, and V' . Therefore if we have:

$$\begin{aligned}
& |\nabla' \rho / \rho| \ll 1, \\
(11) \quad & |\nabla' \lambda / (\lambda+2\mu)| \ll 1, \\
& |\nabla' \mu / (\lambda+2\mu)| \ll 1,
\end{aligned}$$

then all coefficients on both sides of equation (10) except the first term on both sides are $\ll 1$. So neglecting all terms in (10) which are $\ll 1$, we obtain:

$$(12) \quad V'^{-2} \partial^2 \theta' / \partial t'^2 = \nabla'^2 \theta'.$$

Returning (12), which is written in a non-dimensional form, into the dimensional one we get:

$$(13) \quad V^{-2} \partial^2 \theta / \partial t^2 = \nabla^2 \theta,$$

where $V = [(\lambda+2\mu)/\rho]^{1/2}$ is a function of position.

Note that since

$$(14) \quad \nabla' [(\lambda+2\mu)/\rho] = \nabla'(\lambda+2\mu)/\rho - (\lambda+2\mu)\nabla' \rho / \rho^2,$$

we have:

$$(15) \quad \nabla' V^2 / V^2 = 2 \nabla' V / V = \nabla' (\lambda + 2\mu) / (\lambda + 2\mu) - \nabla' \rho / \rho.$$

Using conditions (11) we have:

$$(16) \quad |\nabla' V / V| \ll 1.$$

Returning (16) into the dimensional form and taking the wave length L as the characteristic length we have:

$$(17) \quad |L \nabla V / V| \ll 1.$$

From the above discussion we know that condition (16) is not always equal to condition (11), but if we have conditions (11), we may certainly deduce condition (16), and also deduce condition (17).

As a result the equation of motion (1) can be simplified to the following scalar wave equation:

$$(18) \quad V^{-2} \partial^2 \phi / \partial t^2 = \nabla^2 \phi,$$

assuming the validity of conditions (11).

3. The Ray Equation Approximation of the Wave Equation

In C. B. Officer's book (1974) the conditions under which the wave equation can be approximated by a ray equation are discussed in detail. Let us present his results here.

Suppose that the solution of Eq.(18) has the simple form:

$$(19) \quad \phi = \Phi(x_j) e^{i(\omega t - k_0 \Psi(x_j))}, j= 1,2,3$$

where, $k_0 = \omega/V_0$, V_0 is a reference velocity, and ω is the circular frequency. The necessary relation between Φ and Ψ are obtained after substituting Eq.(19) into Eq.(18) and equating real and imaginary parts:

$$(20) \quad (\nabla \Psi)^2 - (k/k_0)^2 - \nabla^2 \Phi / (k_0^2 \Phi) = 0,$$

$$(21) \quad 1/2 \nabla^2 \Psi + \nabla \Phi \cdot \nabla \Psi / \Phi = 0,$$

where, $k = \omega/V$.

If the following condition is valid

$$(22) \quad |L^2 V''/V| \ll 1,$$

Eq.(20) becomes:

$$(23) \quad (\nabla \Psi)^2 = n^2.$$

That is the so-called eikonal equation, where $n = V_0/V$. In Eq.(22) we have:

$$(24) \quad V'' \approx \partial^2 V / \partial x_1^2.$$

The eikonal equation is a first order differential equation. Its solution

$$(25) \quad \Psi(x_1, x_2, x_3) = \text{constant},$$

represents a surface in a three dimensional space. By choosing the travel time, τ , as a parameter Eq.(25) can be rewritten for different times as

$$(26) \quad \Psi(x_1, x_2, x_3) = \tau.$$

The normals to this wavefront define the direction of wave propagation in any isotropic medium. These are called the rays or ray paths. The ray solution, then is a complete solution to any particular propagation problem within the validity of the approximation of the eikonal equation to the wave equation.

The equation for the normals to the wavefront; i.e. the ray path equation, which will be produced in detail in section 6 is:

$$(27) \quad d(n \, d\vec{r}(x_i)/ds)/ds = \nabla n,$$

where $\vec{r}(x_i)$ is a position vector of the point (x_i) on the ray. The differential ds is an incremental length along the ray path.

We will call (22) the ray approximation condition, as only under this condition we can use the solution to the ray

equation (27) for the description of the kinematic properties of the wave equation (18).

4. The Energy (Amplitude) Equation

We know from the discussion in section (3) that Eq.(20) and Eq.(21) are equivalent to the scalar wave equation (18). From Eq.(20) we obtained the eikonal equation (27) which describes the kinematic characteristics of motion. Here we shall discuss the energy characteristics of motion using Eq.(21).

Choosing the arc length s of ray and q_2 and q_3 as the ray parameters a general curvilinear coordinate system is obtained (see Fig.1). Here q_2 and q_3 are the Gauss coordinates of the curved surface. The unit vectors tangent to the coordinate lines at the point along the ray are:

$$(28) \quad \begin{aligned} \vec{e}_1 &= \partial \vec{r} / \partial s, \\ \vec{e}_i &= \partial \vec{r} / \partial q_i, \end{aligned} \quad (i = 2, 3)$$

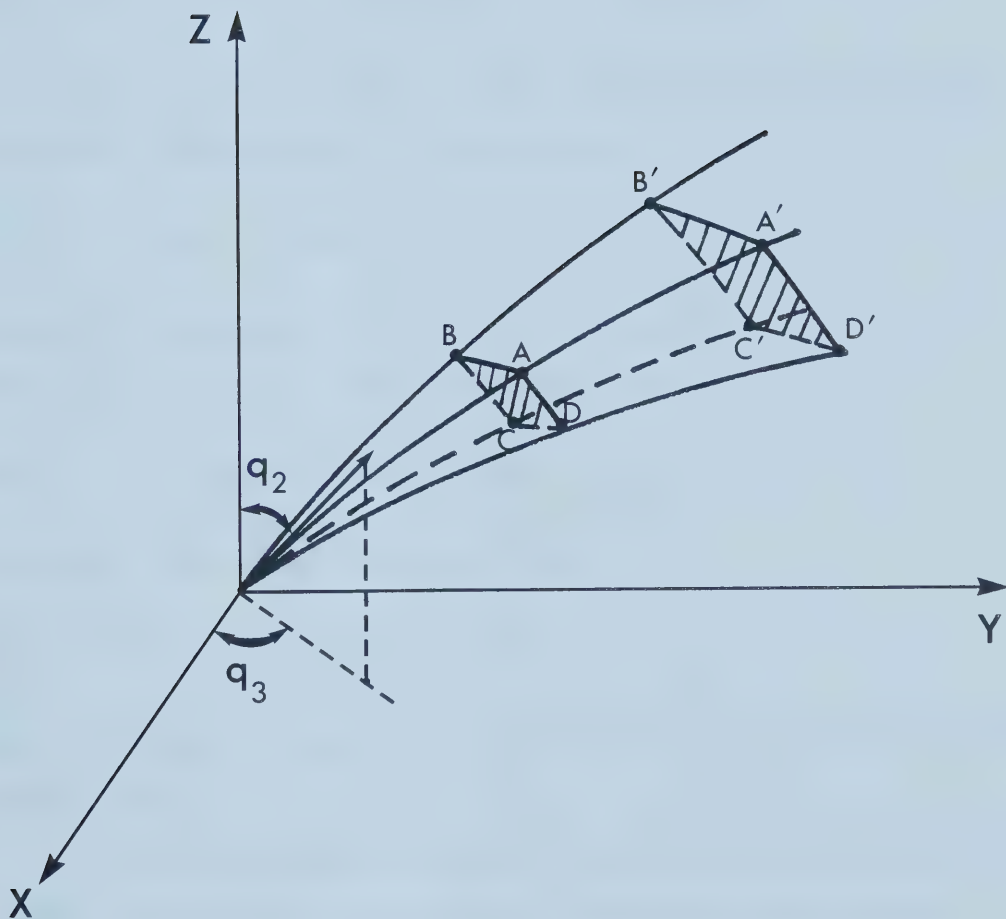
where $\vec{r}$ is the position vector of the point on the ray under consideration.

Lame's coefficients in our curvilinear coordinate system are:

$$(29) \quad \begin{aligned} h_1 &= 1, \\ h_2 &= [(\partial x_1 / \partial q_2)^2 + (\partial x_2 / \partial q_2)^2 + (\partial x_3 / \partial q_2)^2]^{1/2}, \\ h_3 &= [(\partial x_1 / \partial q_3)^2 + (\partial x_2 / \partial q_3)^2 + (\partial x_3 / \partial q_3)^2]^{1/2}. \end{aligned}$$

In the above coordinate system we have:

FIGURE 1. Curvilinear coordinate system on the ray.



$$(30) \quad \nabla \Psi = (\partial \Psi / \partial s, 0, 0),$$

$$(31) \quad \nabla^2 \Psi = (h_2 h_3)^{-1} [\partial (h_2 h_3 \partial \Psi / \partial s) / \partial s],$$

$$(32) \quad \nabla \Phi = (\partial \Phi / \partial s, h_2^{-1} \partial \Phi / \partial q_2, h_3^{-1} \partial \Phi / \partial q_3).$$

As the original scalar wave equation (18) implies the validity of condition (11), we have:

$$(33) \quad |(\lambda + 2\mu)^{-1} \partial (\lambda + 2\mu) / \partial s| \ll 1.$$

Taking $V_0 = 1$ we have $n = 1/V$. By using Eq.(23), (30), (31), (32) and (33) we can rewrite Eq.(21) as:

$$(34) \quad 1/2 \partial \ln[(\lambda + 2\mu) h_2 h_3 / V] / \partial s + \partial \ln \Phi / \partial s = 0.$$

Obviously, the solution of (34) is:

$$(35) \quad \Phi = [(\lambda + 2\mu) h_2 h_3 / V]^{-1/2} C_1,$$

where C_1 is a constant which is associated with the initial conditions.

In Fig.1 we depicted a small area ABCD at point A, and a small area A'B'C'D' at point A'. Denoting the average energy flux due to a simple harmonic wave of frequency ω by E we have:

$$E = 1/2 \omega^2 [\rho \Phi^2 v h_2 h_3 dq_2 dq_3]_A,$$

where A indicates that every quantity has to be taken at point A. So when we put Eq.(35) into the above formula we have

$$(36) \quad E = C_2 \omega^2 dq_2 dq_3,$$

where $C_2 = 2^{-1} C_1^2$. In the same way we can get

$$(37) \quad E' = C_2 \omega^2 dq_2 dq_3,$$

which leads to

$$(38) \quad E = E'.$$

This means that the energy is propagating along the ray tube without any leakage through the sides of the tube. In many practical problems we are frequently asked to find the amplitude of the ray (or equivalently the geometrical spreading) along the ray. We usually choose some parameters having practical geometrical meanings as the parameters q_2 and q_3 . The two takeoff angles α and β serve this purpose quite well. In a generally inhomogeneous medium the two unit vectors $\vec{e}_2$ and $\vec{e}_3$ are tangent to the coordinate lines perpendicular to the ray, but are not usually perpendicular to each other. Nevertheless, the average intensity of the ray at points A and A' are:

$$I = E / (|\vec{e}_2 \times \vec{e}_3| dq_2 dq_3)_A,$$

$$I' = E / (|\vec{e}_2 \times \vec{e}_3| dq_2 dq_3)_{A'}.$$

respectively.

Thus we have

$$(39) \quad I/I' = |\vec{e}_2 \times \vec{e}_3|_{A'} / |\vec{e}_2 \times \vec{e}_3|_A = J(s)/J(s_0).$$

where

$$J(s_0) = |\vec{e}_2 \times \vec{e}_3|_A,$$

$$J(s) = |\vec{e}_2 \times \vec{e}_3|_{A'},$$

Here, s stands for the length of the ray path from some reference point along the ray. Moreover, according to the discussion of average energy in unit time about the simple harmonic wave we may express the wave intensities at A and A' as:

$$I = 1/2 \omega^2 \rho(s_0) \Phi^2(s_0) V(s_0),$$

$$I' = 1/2 \omega^2 \rho(s) \Phi^2(s) V(s),$$

This leads to the well known relation:

$$(40) \quad \Phi(s) = [J(s_0) \rho(s_0) V(s_0) / (J(s) \rho(s) V(s))]^{1/2} \Phi(s_0).$$

which expresses the ray amplitude as a function of the position s along the ray.

Based on the knowledge of differential geometry we have

$$J = (EG - F^2)^{1/2},$$

where:

$$E = \sum_{i=1}^3 (\partial X_i / \partial q_2)^2,$$

$$(41) \quad G = \sum_{i=1}^3 (\partial X_i / \partial q_3)^2,$$

$$F = \sum_{i=1}^3 (\partial^2 X_i / (\partial q_2 \partial q_3)).$$

In practical problems R_1 , the radiation function of the source and R_2 , the receiver function on the earth's surface have to be considered. Thus we can rewrite Eq.(40) as:

$$(42) \quad \Phi(s) = R_1 R_2 [J(s_0)\rho(s_0)V(s_0)/(J(s)\rho(s)V(s))]^{1/2} \Phi(s_0).$$

Note that we have used conditions (11) and (22) during the derivation of Eq.(40) - (42). Hence, the formulae are valid assuming the validity of conditions (11) and (22).

We can see that the energy equation (the amplitude of the ray) requires the same approximating conditions as does the eikonal equation.

5. Discussion on the Relationship of the Approximating Conditions

So far we have used the approximating condition (11), under which the equation of motion (1) leads to the scalar wave equation. We will refer to Eq.(11) as condition (I).

Similar to Eq.(22), under which the scalar wave equation was transformed into the eikonal equation and the amplitude equation, we may write

$$(43) \quad \partial V / \partial x_1 \approx V'.$$

Then Eq.(17) may be rewritten as follows:

$$(44) \quad |L V' / V| \ll 1.$$

which will be called condition (II).

The above mentioned conditions (22), i.e.

$$(22) \quad |L^2 V'' / V| \ll 1.$$

will be called condition (III).

In section 2 we discussed the relationship between condition (I) and condition (II). We showed that (II) follows from (I). Now let us discuss the relationship between condition (II) and condition (III).

In other words, let us investigate the relation between the speed of variation of the elastic parameters of the medium (degree of inhomogeneity) and the minimum frequency (maximum wavelength) of a harmonic source whose seismic wave

field can be decomposed into individual ray contributions. According to B. Officer (1974) and the previous discussion, the inhomogeneity of the medium is always determined by a dimensional length. Officer used L/K as the dimensional length and took $K=1$, where K is a positive constant number and L is one wave length. Let us also take L as our dimensional length.

Therefore, V' and V'' placed in condition (II) and condition (III) may be the first order and the second difference instead of the first and the second derivatives, respectively.

$$(45) \quad |\delta V/V| \ll 1,$$

$$(46) \quad |\delta(\delta V)/V| \ll 1,$$

where δV and $\delta(\delta V)$ is the difference of V and δV in one wave length, respectively.

Obviously if condition (45) is valid, condition (46) is also valid. So we have the following relationship:

$$(47) \quad \text{Condition (I)} \Rightarrow \text{Condition (II)} \Rightarrow \text{Condition (III)}.$$

Relation (47) means that if condition (I) is satisfied we can transform the equation of motion (1) into the scalar wave equation (18). Because at the same time, the ray approximation condition (III) is also valid for any weakly inhomogeneous medium and we can use the eikonal equation (23) and the amplitude equation (42) to study the kinematic and dynamic characteristics of the equation of motion,

assuming the validity of condition (I) without any additional restrictions.

6. The Two Point Seismic Ray Tracing Problem in an Inhomogeneous Medium

We will consider here the case of an inhomogeneous isotropic medium. Let us assume that there is a velocity of propagation, $V(x,y,z)$, at each point $P(x,y,z)$ of the inhomogeneous isotropic medium. We further assume that $V(x,y,z)$ possesses continuous first partial derivatives. The following shows the derivation of the ray equation (27), based on Fermat's principle, in detail.

6.1 The Derivation of the Ray Path Equation

Let us draw an arbitrary curve which connects points $A(x_0, y_0, z_0)$ and $B(x_1, y_1, z_1)$. A particle moving from A to B along the path Γ has the assigned velocity $V(x,y,z)$ at each (x,y,z) position, and moves from A to B in time T given by

$$(48) \quad T = \int_A^B ds/V(x,y,z),$$

where ds is an element along the ray path Γ .

According to Fermat's principle the time T taken by the seismic ray to travel from A to B has a stationary value.

Let us introduce a new parameter ξ to characterize the ray path and write

$$(49) \quad \begin{aligned} x &= x(\xi), \\ y &= y(\xi), \\ z &= z(\xi). \end{aligned}$$

An element of the ray path ds is then written as:

$$(50) \quad ds = (dx^2 + dy^2 + dz^2)^{1/2} = (\dot{x}^2 + \dot{y}^2 + \dot{z}^2)^{1/2} d\xi,$$

where:

$$\dot{x} = dx/d\xi, \quad \dot{y} = dy/d\xi, \quad \dot{z} = dz/d\xi.$$

Eq.(48) becomes

$$(51) \quad T = \int_{\xi_1}^{\xi_2} (\dot{x}^2 + \dot{y}^2 + \dot{z}^2)^{1/2} / V(x, y, z) d\xi,$$

where $\xi = \xi_1$ at A, and $\xi = \xi_2$ at B. This equation is rewritten as follows:

$$(52a) \quad T = \int_{\xi_1}^{\xi_2} W d\xi,$$

$$(52b) \quad W \equiv [\dot{x}(\xi)^2 + \dot{y}(\xi)^2 + \dot{z}(\xi)^2]^{1/2} / V(x(\xi), y(\xi), z(\xi)).$$

Obviously $W(\xi, x(\xi), y(\xi), z(\xi), \dot{x}(\xi), \dot{y}(\xi), \dot{z}(\xi))$ is a function of ξ and has a stationary value. The Euler equation:

$$d(\partial W / \partial \dot{Y}') / d\xi - \partial W / \partial Y = 0$$

requires

$$(53) \quad \begin{aligned} d(\partial W / \partial \dot{x}) / d\xi - \partial W / \partial x &= 0, \\ d(\partial W / \partial \dot{y}) / d\xi - \partial W / \partial y &= 0, \\ d(\partial W / \partial \dot{z}) / d\xi - \partial W / \partial z &= 0. \end{aligned}$$

Substituting (52b) into (53) we have:

$$\begin{aligned}
 & d[V^{-1}\dot{x}/(\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}]/d\xi - (\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}\partial(1/V)/\partial x = 0, \\
 (54) \quad & d[V^{-1}\dot{y}/(\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}]/d\xi - (\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}\partial(1/V)/\partial y = 0, \\
 & d[V^{-1}\dot{z}/(\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}]/d\xi - (\dot{x}^2+\dot{y}^2+\dot{z}^2)^{1/2}\partial(1/V)/\partial z = 0.
 \end{aligned}$$

The parameter ξ along the ray path is still arbitrary, so if we set $\xi=s$, the arc length of Γ , then in order to satisfy Eq.(50) we must have:

$$(55) \quad (dx/ds)^2 + (dy/ds)^2 + (dz/ds)^2 = 1,$$

along Γ . Hence Eq.(54) may be simplified to:

$$\begin{aligned}
 & d(V^{-1}dx/ds)/ds - \partial(1/V)/\partial x = 0, \\
 (56) \quad & d(V^{-1}dy/ds)/ds - \partial(1/V)/\partial y = 0, \\
 & d(V^{-1}dz/ds)/ds - \partial(1/V)/\partial z = 0.
 \end{aligned}$$

If we take $\vec{r}(s)$ as the space curve defining the ray path joining the two points A and B, then Eq.(56) amounts to:

$$(57) \quad d(V^{-1}d\vec{r}/ds)/ds = \nabla(1/V).$$

After choosing $V_0=1$ in $n=V_0/V$, Eq.(27) is identical to (57). We shall use the ray equation (57) in the following form:

$$(57') \quad d(u(\vec{r})d\vec{r}/ds)/ds = \nabla(u(\vec{r})),$$

where $\vec{r} = (x, y, z)$ and $u \equiv 1/V(\vec{r})$.

6.2 Two Point Seismic Ray Tracing Equation in 3-D

Inhomogeneous Media

Let us first recast equation (57') as:

$$(58) \quad d^2\vec{r}/ds^2 = v(\vec{r})[\nabla u(\vec{r}) - G(\vec{r})d\vec{r}/ds],$$

where,

$$G(\vec{r}) = (\partial u/\partial x)(dx/ds) + (\partial u/\partial y)(dy/ds) + (\partial u/\partial z)(dz/ds).$$

Writing (58) in detail leads to the following formulation for the two point ray tracing problem:

$$(59) \quad \begin{aligned} d^2x/ds^2 &= v(\vec{r})[\partial u/\partial x - G(\vec{r})dx/ds], \\ d^2y/ds^2 &= v(\vec{r})[\partial u/\partial y - G(\vec{r})dy/ds], \\ d^2z/ds^2 &= v(\vec{r})[\partial u/\partial z - G(\vec{r})dz/ds]. \end{aligned}$$

If S is the total length of the ray between A and B, then the boundary conditions are:

$$(60) \quad \begin{aligned} x(0) &= x_0, \quad y(0) = y_0, \quad z(0) = z_0; \\ x(S) &= x_1, \quad y(S) = y_1, \quad z(S) = z_1. \end{aligned}$$

The main advantage of using an arc length parameterization lies in the fact that the functions $x(s)$, $y(s)$ and $z(s)$ are single valued, even when rays intersect.

It is convenient to express Eq.(59) as a system of first order nonlinear differential equations. It is known that second or higher order systems of ordinary differential equations can be easily reduced to first order systems by

introducing auxiliary variables. For instance Eq.(59) reduces to:

$$\begin{aligned}
 dY_1/ds &= Y_2, \\
 dY_2/ds &= V[\partial u/\partial x - G(\vec{Y})Y_2], \\
 (61) \quad dY_3/ds &= Y_4, & 0 \leq s \leq S \\
 dY_4/ds &= V[\partial u/\partial y - G(\vec{Y})Y_4], \\
 dY_5/ds &= Y_6, \\
 dY_6/ds &= V[\partial u/\partial z - G(\vec{Y})Y_6],
 \end{aligned}$$

where,

$$(62) \quad G(\vec{Y}) = u_x Y_2 + u_y Y_4 + u_z Y_6,$$

and the vector $\vec{Y} \equiv (x, dx/ds, y, dy/ds, z, dz/ds)^T$.

Since in many of these calculations the travel time $T = \int u ds$ is of interest, a new variable, Y_7 , may be introduced to represent the local travel time. This necessitates the introduction of an additional differential equation

$$dY_7/ds = u,$$

with the initial condition

$$Y_7(0) = 0.$$

In this manner the total travel time $T = Y_7(S)$ is computed automatically with the corresponding ray path.

In order to determine the unknown parameter S we change variable s to $\tau = s/S$ and introduce a trivial differential equation $S' = 0$, where now the prime (') will denote differentiation with respect to τ . Calling $Y_8 = S$, we obtain

the final set of equations:

$$\begin{aligned}
 Y_1' &= Y_8 Y_2, \\
 Y_2' &= Y_8 V[\partial u / \partial x - G(Y) Y_2], \\
 Y_3' &= Y_8 Y_4, \\
 (63) \quad Y_4' &= Y_8 V[\partial u / \partial y - G(Y) Y_4], & 0 \leq \tau \leq 1 \\
 Y_5' &= Y_8 Y_6, \\
 Y_6' &= Y_8 V[\partial u / \partial z - G(Y) Y_6], \\
 Y_7' &= Y_8 u, \\
 Y_8' &= 0.
 \end{aligned}$$

After introducing τ , a particle moving from $A(x_0, y_0, z_0)$ to $B(x_1, y_1, z_1)$ will produce the change of τ from 0 to 1. So the corresponding boundary conditions to Eq.(63) are:

$$\begin{aligned}
 Y_1(0) &= x_0, \\
 Y_3(0) &= y_0, \\
 Y_5(0) &= z_0, \\
 (64) \quad Y_1(1) &= x_1, \\
 Y_3(1) &= y_1, \\
 Y_5(1) &= z_1, \\
 Y_7(0) &= 0.
 \end{aligned}$$

The eighth boundary condition which is needed to completely define the system can be obtained from Eq.(55) which can be expressed in terms of Y as

$$(65) \quad Y_2(0)^2 + Y_4(0)^2 + Y_6(0)^2 = 1,$$

where

$$(66) \quad \vec{Y} \equiv (x, dx/ds, y, dy/ds, z, dz/ds, t, S)^T.$$

Solving Eq.(63) with the boundary conditions Eq.(64) and (65) will give the solution to the three dimensional, two point ray tracing problem.

6.3 Two Point Seismic Ray Tracing Equation in 2-D

Inhomogeneous Media

In the case of a two dimensional inhomogeneous medium we may only consider the two point ray tracing problem in the x-z plane. Eq.(56) becomes:

$$(67) \quad \begin{aligned} d(u(x,z) dx/ds)/ds - \partial u(x,z)/\partial x &= 0, \\ d(u(x,z) dz/ds)/ds - \partial u(x,z)/\partial z &= 0. \end{aligned}$$

According to Eq.(55) we have:

$$(68) \quad (dx/ds)^2 + (dz/ds)^2 = 1.$$

Let us introduce a new parameter ψ defined by

$$(69) \quad \begin{aligned} dx/ds &= \cos\psi, \\ dz/ds &= \sin\psi. \end{aligned}$$

We find that

$$\begin{aligned} d^2x/ds^2 &= -(dz/ds)(d\psi/ds), \\ d^2z/ds^2 &= (dx/ds)(d\psi/ds), \end{aligned}$$

yielding

$$(70) \quad (d^2x/ds^2)(dz/ds) - (d^2z/ds^2)(dx/ds) = -d\psi/ds.$$

Expanding equations (67) and multiplying respectively by dz/ds and dx/ds and subtracting, we obtain

$$u(d^2x/ds^2 dz/ds - d^2z/ds^2 dx/ds) + u_z dx/ds - u_x dz/ds = 0.$$

Using Eq.(70) we have

$$(71) \quad d\psi/ds = V(u_z dx/ds - u_x dz/ds).$$

Introducing the variables

$$(72) \quad Y \equiv (Y_1, Y_2, Y_3, Y_4, Y_5)^T = (x, z, \psi, S, t)^T,$$

and changing in variables s to $\tau=s/S$, the resulting, first order system is

$$\begin{aligned} Y_1' &= Y_4 \cos Y_3, \\ Y_2' &= Y_4 \sin Y_3, \\ (73) \quad Y_3' &= Y_4 V[u_z \cos Y_3 - u_x \sin Y_3], \quad 0 \leq \tau \leq 1 \\ Y_4' &= 0, \\ Y_5' &= Y_4 u, \end{aligned}$$

where now the prime (') denotes differentiation with respect to τ . Obviously the corresponding boundary conditions are:

$$\begin{aligned} Y_1(0) &= x_0, \\ Y_2(0) &= z_0, \\ (74) \quad Y_1(1) &= x_1, \\ Y_2(1) &= z_1, \\ Y_5(0) &= 0. \end{aligned}$$

Solving Eq.(73) with the boundary conditions (74) will give us the solution of the two point ray tracing problem in a two dimensional inhomogeneous medium.

6.4 An Alternative Formulation of the Two Point Ray Tracing Problem in 2-D Media

Using x-z coordinates and taking the ray path as a function of the x-coordinate, the differential ray length ds and the travel time T are, respectively:

$$(75) \quad ds = [1+(dz/dx)^2]^{1/2} dx,$$

$$(76) \quad T = \int_{x_0}^{x_1} [1+(dz/dx)^2]^{1/2}/V(x,z) dx.$$

According to Euler's equations (53) we have:

$$W = [1+(dz/dx)^2]^{1/2}/V(x,z),$$

$$d[\partial W/\partial \dot{z}]/dx - \partial W/\partial z = 0,$$

and finally

$$(77) \quad d^2z/dx^2 = V^{-1}[1+(dz/dx)^2][(\partial V/\partial x)(dz/dx) - \partial V/\partial z].$$

Introducing two new auxiliary parameters Z_1 and Z_2 as follows,

$$(78) \quad Z_1 \equiv z,$$

$$Z_2 \equiv dZ_1/dx = dz/dx,$$

Eq.(77) becomes

$$dZ_1/dx = Z_2,$$

$$(79) \quad \frac{dz_2}{dx} = V^{-1}(1+z_2^2)[(\partial V/\partial x)z_2 - \partial V/\partial z_1], \quad x_0 \leq x \leq x_1$$

with the relative boundary conditions:

$$(80) \quad \begin{aligned} z_1(x_0) &= z_0, \\ z_1(x_1) &= z_1. \end{aligned}$$

Solving Eq.(79) with (80) will give the solution of the ray path between two end points $A(x_0, z_0)$ and $B(x_1, z_1)$ in a two dimensional medium.

6.5 Treatment of Interfaces in the Medium

From the above discussion we can see that the seismic ray tracing problem has been reduced to solving a set of first order differential equations of the form

$$(81) \quad \vec{Y}' = \vec{F}(\tau, \vec{Y}(\tau)), \quad \tau \in [0, 1]$$

with the boundary conditions formally written as

$$\vec{g}[\vec{Y}(0), \vec{Y}(1)] = 0,$$

where $\vec{Y}$, $\vec{F}$, $\vec{g}$ are all d - dimensional vector functions, and defined in a R -dimensional space. $\vec{F}$ is a nonlinear function.

In ray tracing problems, as in many other applications, it is possible for problems (64) or (73) to have jump discontinuities within the region of interest. These discontinuities appear in the function $\vec{F}(\tau, \vec{Y})$. V. Pereyra et al.(1980) have discussed this problem. If $\vec{F}$ is discontinuous with respect to the independent variable at a known location δ_1 , $0 < \delta_1 < 1$, then the solution $\vec{Y}^*(\tau)$ or some of its

derivatives may be discontinuous at δ_1 , as well. If $\vec{F}$ is discontinuous with respect to $\vec{Y}$ then the location of the possible discontinuities in $\vec{Y}^*(\tau)$ or its derivatives are not known a priori.

Assume for simplicity that $\vec{F}$ becomes discontinuous when the ray or trajectory $\vec{Y}^*(\tau)$, traverses a given surface $\Phi(\tau, \vec{Y})=0$ in $(0,1) \in \mathbb{R}^d$ and that this occurs only once for $\tau \in (0,1)$. Thus the condition for $\vec{Y}^*(\tau)$ to lie on this surface is that $\Phi(\tau, \vec{Y}^*(\tau))=0$. We will call δ_1 that value of τ for which the ray touches the interface.

Assuming that the desired solution has a known behavior across the discontinuity, i.e. obeys a jump discontinuity condition of the form $\vec{D}[\vec{Y}^*(\delta_1, -), \vec{Y}^*(\delta_1, +)] = 0$ where $\vec{D}$ is d -dimensional and " $\pm$ " represent the value on the right and left of the discontinuity, this problem can be reduced to the form (64) or (73) for some data. Any number of interfaces of this type can be treated simultaneously in the same way, provided one knows the order in which they are traversed and in which region the ray is situated, i.e. which velocity function applies. It is obvious that one can as easily consider both reflected and refracted rays, as well as a change from P to S waves and vice versa.

The idea of the reduction is to consider system (73) independently in two regions $[0, \delta_1]$, $[\delta_1, 1]$, and use the interface condition $\vec{D}=0$, to couple these two problems. In each one of the intervals the problem is smooth. In the case of known crossing time δ_1 , a system twice the size of the

original one is considered (i.e. one subsystem for each subinterval), adding the interface conditions and the necessary d additional boundary conditions.

The case of an unknown crossing time requires the introduction of an auxiliary dependent variable $\delta_1(\tau)$, which satisfies the trivial first order differential equation $\delta_1'(\tau)=0$. The additional boundary condition is now given by the definition of the interface and the requirement that $\vec{Y}^*$ lie on it for $\tau=\delta_1$:

$$\Phi[\delta_1, \vec{Y}^*(\delta_1)] = 0.$$

If we call $[0, \delta_1]$, $[\delta_1, 1]$ as regions (1) and (2), respectively, map each interval time into $[0, 1]$ and call $\vec{Y}_{(1)}$ and $\vec{Y}_{(2)}$ the solution on these subintervals, we have an augmented system of the form:

$$\begin{aligned} \vec{Y}_{(1)}' &= \delta_1 \vec{F}(\tau, \vec{Y}_{(1)}), \\ (82) \quad \vec{Y}_{(2)}' &= (1-\delta_1) \vec{F}(\tau, \vec{Y}_{(2)}), \quad \tau \in [0, 1] \\ \delta_1' &= 0, \end{aligned}$$

with the boundary conditions:

$$\begin{aligned} \vec{g}[\vec{Y}_{(1)}(0), \vec{Y}_{(2)}(1)] &= 0, \\ (83) \quad \vec{D}[\vec{Y}_{(1)}(1), \vec{Y}_{(2)}(0)] &= 0, \\ \Phi[\delta_1(1), \vec{Y}_{(1)}(1)] &= 0. \end{aligned}$$

Apparently the size of the problem has been greatly increased from d equations to $(2*d+1)$ equations, or in the case of K interfaces to $(K+1)*d+K$ equations. However, the

real dimension of the problem is $N \cdot E$, E being number of equations and N the number of mesh points.

6.6 A Simple Example for the Treatment of the Interfaces in a 2-D Medium

The theory just developed can now be applied to equation (73). For simplicity, we will start with the case which has only one discontinuity in the velocity field, i.e. when the ray crosses the curve. let

$$(84) \quad \vec{Y} \equiv (Y_1, Y_2, Y_3, Y_4, Y_5, Y_6, Y_7, Y_8, Y_9, Y_{10})^T, \\ \equiv (x_{(1)}, z_{(1)}, \psi_{(1)}, S_{(1)}, t_{(1)}, x_{(2)}, z_{(2)}, \psi_{(2)}, S_{(2)}, t_{(2)})^T.$$

From Eq.(73) and the theory described in section 6.5 we may have the following set of the differential equations:

$$(85) \quad \begin{aligned} Y_1' &= Y_4 \cos Y_3, \\ Y_2' &= Y_4 \sin Y_3, \\ Y_3' &= Y_4 V_{(1)} (u_{(1)z} \cos Y_3 - u_{(1)x} \sin Y_3), \\ Y_4' &= 0, \\ Y_5' &= Y_4 u_{(1)}, \\ Y_6' &= Y_9 \cos Y_8, \\ Y_7' &= Y_9 \sin Y_8, \\ Y_8' &= Y_9 V_{(2)} (u_{(2)x} \cos Y_8 - u_{(2)z} \sin Y_8), \\ Y_9' &= 0, \\ Y_{10}' &= Y_9 u_{(2)}. \end{aligned} \quad 0 \leq \tau \leq 1$$

Two new conditions will be given, by Snell's law, which the ray must obey when impinging on an interface separating

two different elastic media, and by the fact that the incident point C has to be on the curve $\Phi(C)=0$ is reached at $\tau=\delta_1$.

Setting $\psi_1(1)$ as the angle of incidence, $\psi_2(0)$ the angle of refraction, and letting θ_1 , θ_2 be the angles between the ray direction and the normal direction of the curve, respectively in region (1) and in region (2), and letting η_1 , η_2 be the angles between the normal direction and the x axis, respectively in region (1) or (2), we have:

$$\theta_1 = \psi_1(1) - \eta_1.$$

Because of

$$\cos\eta_1 = \Phi_x / (\Phi_x^2 + \Phi_z^2)^{1/2},$$

$$\sin\eta_1 = \Phi_z / (\Phi_x^2 + \Phi_z^2)^{1/2},$$

we have

$$(86) \quad \begin{aligned} \sin\theta_1 &= [\Phi_x \sin\psi_1(1) - \Phi_z \cos\psi_1(1)] / (\Phi_x^2 + \Phi_z^2)^{1/2}, \\ \sin\theta_2 &= [\Phi_x \sin\psi_2(0) - \Phi_z \cos\psi_2(0)] / (\Phi_x^2 + \Phi_z^2)^{1/2}. \end{aligned}$$

Employing Snell's law we have:

$$(87) \quad \begin{aligned} V_{(2)}(c) [\Phi_x \sin\psi_1(1) - \Phi_z \cos\psi_1(1)] &= \\ V_{(1)}(c) [\Phi_x \sin\psi_2(0) - \Phi_z \cos\psi_2(0)]. \end{aligned}$$

Then we have the following boundary conditions subject to Eq.(85):

$$Y_1(0) - x_0 = 0,$$

$$Y_2(0) - z_0 = 0,$$

$$\begin{aligned}
 &Y_5(0) = 0, \\
 &Y_6(1) - x_1 = 0, \\
 &Y_7(1) - z_1 = 0, \\
 (88) \quad &Y_{10}(0) = 0, \\
 &Y_6(0) - Y_1(1) = 0, \\
 &\Phi(Y_1(1), Y_2(1)) = 0, \\
 &\Phi(Y_6(0), Y_7(0)) = 0, \\
 &V_{(2)}(Y_6(0), Y_7(0)) \cdot \\
 &\quad \cdot [\Phi_x(Y_1(1), Y_2(1)) \sin Y_3(1) - \Phi_z(Y_1(1), Y_2(1)) \cos Y_3(1)] \\
 &= V_{(1)}(Y_1(1), Y_2(1)) \cdot \\
 &\quad \cdot [\Phi_x(Y_6(0), Y_7(0)) \sin Y_8(0) - \Phi_z(Y_6(0), Y_7(0)) \cos Y_8(0)].
 \end{aligned}$$

7. A Numerical Method for Solving Ray Equations

A numerical method used to solve equations of the form Eq.(81) (of which the ray tracing equations are examples) is based on an adaptive difference solver for nonlinear two point boundary problems described by two mathematicians V. Pereyra and H. B. Keller and a seismologist W. H. K. Lee (1980). The details of the mathematics of this solver can be found by interested readers in the following set of mathematical papers: H. B. Keller (1968,1969,1974); V. Pereyra (1967,1968,1973); M. Lentini and V. Pereyra (1974, 1975, 1977); A. Bjorck and V. Pereyra (1970). Let us simply outline the method. Later we will refer to this solver as Pereyra's method.

Pereyra's method is based on a simple finite difference approximation to $d\vec{Y}/d\tau$ on a mesh with $(n+1)$ points in the interval $[0,1]$. Consider a mesh π of points $\{\tau_j\}$, $j = 1,2,\dots,n+1$, satisfying:

$$(89) \quad 0 = \tau_1 < \tau_2 < \dots < \tau_{n+1} = 1.$$

In order to change the nonlinear differential equations into nonlinear difference equations (nonlinear algebraic equations), the basic finite difference approximation to Eq.(81) is considered for the trapezoidal rule:

$$(90) \quad \begin{aligned} \vec{\Phi}_{\pi}(\vec{W})_j &\equiv (\vec{W}_{j+1} - \vec{W}_j)/h_j - 1/2[\vec{F}(\tau_{j+1}, \vec{W}_{j+1}) + \vec{F}(\tau_j, \vec{W}_j)] \\ &= 0, \end{aligned} \quad j = 1,2,\dots,n$$

$$\vec{g}(\vec{w}_1, \vec{w}_{n+1}) = 0.$$

Equations (90) form a system of $(n+1)d$ nonlinear algebraic equations. Using Newton interaction we may change the nonlinear problem (90) into a linear problem, i.e.

$$(91) \quad \vec{w}^{(k+1)} = \vec{w}^{(k)} - [\vec{\Phi}_\pi'(\vec{w}^{(k)})]^{-1} \vec{\Phi}_\pi(\vec{w}^{(k)}),$$

where $\vec{\Phi}_\pi'$ is the Jacobian matrix of $\vec{\Phi}_\pi$:

$$(91) \quad \vec{\Phi}_\pi'(\vec{w}) = \begin{bmatrix} \vec{G}_1, & 0, & \dots\dots\dots, & 0, & \vec{G}_{n+1} \\ \vec{S}_1, & \vec{R}_1, & 0, & \dots\dots\dots, & 0 \\ 0, & \vec{S}_2, & \vec{R}_2, & \dots\dots\dots, & 0 \\ \vdots & & \ddots & \ddots & \vdots \\ 0, & \dots\dots\dots, & 0, & \vec{S}_n, & \vec{R}_n \end{bmatrix}$$

where:

$$\vec{G}_1 = \partial \vec{g} / \partial w_{1i},$$

$$\vec{G}_{n+1} = \partial \vec{g} / \partial w_{n+1i};$$

$$\vec{S}_j = -I/h_j - 1/2 \partial \vec{F}(\tau_j, \vec{w}_j) / \partial w_{ji},$$

$$\vec{R}_j = I/h_j - 1/2 \partial \vec{F}(\tau_{j+1}, \vec{w}_{j+1}) / \partial w_{j+1i},$$

$$i = 1, 2, \dots\dots\dots, d. \quad j = 2, \dots\dots\dots, J.$$

In order to obtain the more accurate approximation we introduce the operators $\vec{S}_\pi^{(k)}$ which are explained in detail in the papers (Pereyra, 1967, 1968; Keller, 1974). Letting $\vec{S}_\pi^{(0)}(u^{(-1)}) \equiv 0$ and solving successively for $k = 0, 1, 2, \dots\dots\dots$

$$(93) \quad \vec{\Phi}_\pi(\vec{w}^{(k)}) = \vec{S}_\pi^{(k)}(\vec{w}^{(k-1)}),$$

$$\vec{g}(\vec{w}_1, \vec{w}_{n+1}) = 0.$$

After introducing the concept of an asymptotically equidistributing mesh (see Lentini et al, 1977; Pereyra and Sewell, 1975) we can use a nonuniform mesh to increase the calculation efficiency.

Based on the above idea mentioned, I have programmed Pereyra's method to solve two point ray tracing problems like Eq.(73), (79) and (85). Let us first show one of the features of this method, consisting of an automatic procedure for choosing a nonuniform mesh using the following simple example.

Example: Consider the first order scalar equation

$$(94) \quad \begin{aligned} Y' &= -100 Y(t), \\ Y(0) &= 1. \end{aligned} \quad t \in [0, 1]$$

whose solution is

$$Y(t) = e^{-100t}.$$

There is a boundary layer of width $1/100$ at $t=0$ (see Fig.2). Using my program, for the accuracy standard of $TOL=10^{-3}$, the total number of mesh points needed was $N=28$. The calculation results with the corresponding analytical solution at the same points is listed in Table 1. We can see of the solution that the mesh points are close together in the region of fast parameter variation ($0.0 - 0.083$), where 11 points are used; while in the open region, ($0.083 - 1.0$), only 17 mesh points were needed. If we use the uniform mesh for the same

accuracy, we have to use a total of $N=129$ points.

The variable mesh technique allows the program to "adapt itself" to the particular problem at hand, and thus it produces more detailed computations where they are needed, as in the case of highly curved rays.

FIGURE 2. Example of the nonuniform mesh technique.

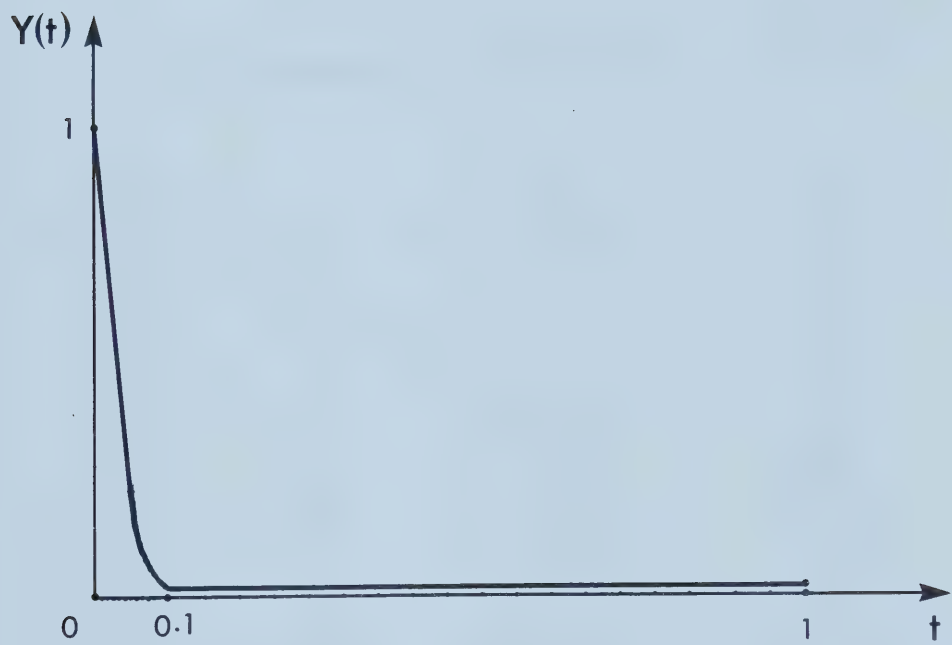


Table 1. Solutions of the Example about the Nonuniform Mesh

N	t	Increment	Analytic Solution	Numerical Solution
1	0.000		1.000	1.000
2	0.007	0.007	0.499	0.499
3	0.014	0.007	0.249	0.250
4	0.021	0.007	0.125	0.125
5	0.028	0.007	0.062	0.063
6	0.035	0.007	0.031	0.031
7	0.042	0.007	0.016	0.016
8	0.049	0.007	0.008	0.008
9	0.056	0.007	0.004	0.004
10	0.063	0.007	0.002	0.002
11	0.083	0.021	0.000	0.000
12	0.104	0.021	0.000	0.000
13	0.125	0.021	0.000	-0.000
14	0.156	0.031	0.000	0.000
15	0.188	0.031	0.000	-0.000
16	0.250	0.063	0.000	0.000
17	0.313	0.063	0.000	-0.000
18	0.375	0.063	0.000	0.000
19	0.438	0.063	0.000	-0.000
20	0.500	0.063	0.000	0.000
21	0.563	0.063	0.000	-0.000
22	0.625	0.063	0.000	0.000
23	0.688	0.063	0.000	-0.000
24	0.750	0.063	0.000	0.000
25	0.813	0.063	0.000	-0.000
26	0.875	0.063	0.000	0.000
27	0.938	0.063	0.000	-0.000
28	1.000	0.063	0.000	0.000

8. Computer Programs for Two Point Seismic Ray Tracing and Some Examples

As we have shown in sections 6 and 7, the mathematics involved in two point seismic ray tracing problems can become rather complex. Therefore, we must consider how to verify that our computer programs are correct. Traditionally, one tests against some simple cases where analytic solutions are known. It is obvious that this is necessary but not sufficient, as in most cases analytic solutions are known for only trivial cases. If two different programs based on two different methods are written to solve the same problem, then we will have confidence in our programs if they give essentially the same answers.

We have written three computer programs implementing Pereyra's method ZH1 solving Eq.(79) with (80), ZH2 solving Eq.(73) with (74) and ZH3 solving Eq.(85) with (88). These solutions were checked against the programs of W. H. K. Lee ,SRT1 and SRT2, which were based on Wesson's and Lentini's methods, respectively. Thus we have been able to verify that our programs are coded correctly and simultaneously estimate the accuracy and efficiency of Pereyra's method.

8.1 Test Case No.1

In the first case, we assume $V(x,z) = 4 + 0.2 z$ (km/sec). i.e. velocity is a linear function of only the z-coordinate. We select the two end points as (0, 10) and

(100, 0) (km) to test a ray path similar to the refracted wave case for vertically inhomogeneous model. The analytic solution is a circular ray path with its center at (47.5, -20.0) (km). The initial guess is a straight line between the two end points (see Fig.3). The computed solution from program ZH1 together with the analytic solution is shown in Table 3. In Table 2, the computed solutions from the programs SRT1 and SRT2 (Lee's) together with the analytic solution are shown. The maximum deviations of the SRT1 and SRT2 solutions from the analytic solution are about 0.16 km and 0.00028 km respectively; whereas the maximum deviation of the ZH1 solution from the analytic one is about 0.00003 km. So the accuracy of program ZH1 is better than that of programs SRT1 and SRT2.

FIGURE 3. Geometry for case No.1.

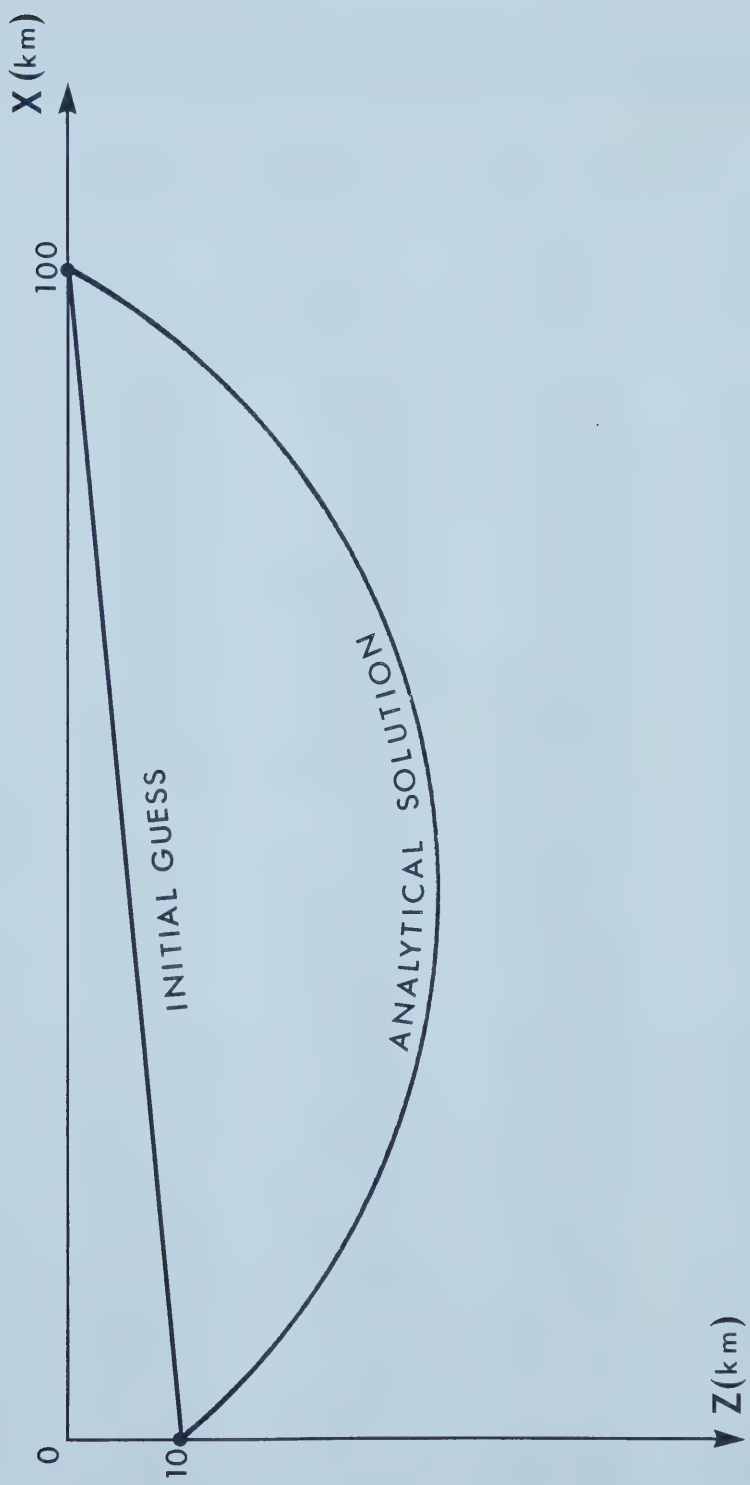


Table 2. Comparison of Analytic, SRT1 and SRT2 Solutions for Case No.1

X	INITIAL GUESS	ANALYTIC SOLUTION	SRT1 SOLUTION	SRT2 SOLUTION	$Z_A - Z_1$	$Z_A - Z_2$
(KM)	Z_0 (KM)	Z_A (KM)	Z_1 (KM)	Z_2 (KM)	(KM)	(KM)
0.00000	10.00000	10.00000	10.00000	10.00000	0.00000	0.00000
1.00000	9.90000	11.52776	11.51953	11.52776	0.00823	-0.00000
2.00000	9.80000	12.95451	12.94046	12.95450	0.01405	0.00001
3.00000	9.70000	14.29285	14.27446	14.29285	0.01839	-0.00000
4.00000	9.60000	15.55276	15.53099	15.55277	0.02178	-0.00000
5.00000	9.50000	16.74234	16.71777	16.74234	0.02457	0.00000
6.00000	9.40000	17.86818	17.84128	17.86918	0.02691	0.00000
7.00000	9.30000	18.93584	18.91691	18.93584	0.02893	0.00000
8.00000	9.20000	19.94997	19.91925	19.94997	0.03072	0.00000
9.00000	9.10000	20.91454	20.88222	20.91454	0.03232	0.00000
10.00000	9.00000	21.83299	21.79922	21.83299	0.03377	0.00000
11.00000	8.90000	22.70831	22.67319	22.70831	0.03513	0.00000
12.00000	8.80000	23.54308	23.50671	23.54308	0.03636	0.00000
13.00000	8.70000	24.33958	24.30206	24.33958	0.03752	0.00000
14.00000	8.60000	25.09998	25.06123	25.09998	0.03865	0.00000
15.00000	8.50000	25.82574	25.78606	25.82574	0.03969	0.00000
16.00000	8.40000	26.51980	26.47809	26.51861	0.04071	-0.00000
17.00000	8.30000	27.18050	27.13881	27.18050	0.04169	0.00000
18.00000	8.20000	27.81212	27.76950	27.81213	0.04262	-0.00000
19.00000	8.10000	28.41487	28.37134	28.41487	0.04353	0.00000
20.00000	8.00000	28.98979	28.94539	28.98979	0.04440	0.00000
21.00000	7.90000	29.53786	29.49260	29.53786	0.04526	0.00000
22.00000	7.80000	30.05995	30.01384	30.05997	0.04611	-0.00000
23.00000	7.70000	30.55698	30.50996	30.55688	0.04692	0.00000
24.00000	7.60000	31.02939	30.98166	31.02940	0.04773	-0.00000
25.00000	7.50000	31.47815	31.42961	31.47815	0.04854	0.00000
26.00000	7.40000	31.90375	31.85443	31.90375	0.04932	0.00000
27.00000	7.30000	32.30678	32.25670	32.30678	0.05008	0.00000
28.00000	7.20000	32.68776	32.63690	32.68774	0.05086	0.00000
29.00000	7.10000	33.04713	32.99553	33.04713	0.05161	0.00000
30.00000	7.00000	33.38539	33.33301	33.38539	0.05238	0.00000
31.00000	6.90000	33.70288	33.64977	33.70288	0.05312	0.00000
32.00000	6.80000	34.00000	33.94612	33.99998	0.05388	0.00000
33.00000	6.70000	34.27705	34.22244	34.27705	0.05461	0.00000
34.00000	6.60000	34.53438	34.47902	34.53439	0.05536	-0.00000
35.00000	6.50000	34.77225	34.71613	34.77225	0.05612	0.00000
36.00000	6.40000	34.99091	34.93404	34.99091	0.05687	0.00000
37.00000	6.30000	35.19057	35.13295	35.19057	0.05762	0.00000
38.00000	6.20000	35.37146	35.31308	35.37146	0.05838	0.00000
39.00000	6.10000	35.53377	35.47462	35.53375	0.05914	0.00000
40.00000	6.00000	35.67764	35.61771	35.67763	0.05994	0.00000
41.00000	5.90000	35.80322	35.74249	35.80322	0.06073	0.00000
42.00000	5.80000	35.91063	35.84912	35.91063	0.06151	0.00000
43.00000	5.70000	36.00000	35.93767	35.99998	0.06233	0.00000
44.00000	5.60000	36.07138	36.00824	36.07137	0.06314	0.00000
45.00000	5.50000	36.12485	36.06090	36.12485	0.06395	0.00000
46.00000	5.40000	36.16048	36.09569	36.16048	0.06479	0.00000
47.00000	5.30000	36.17828	36.11254	36.17827	0.06564	0.00000
48.00000	5.20000	36.17828	36.11180	36.17827	0.06644	0.00000
49.00000	5.10000	36.16048	36.09312	36.16048	0.06735	0.00000
50.00000	5.00000	36.12485	36.05664	36.12485	0.06821	0.00000
51.00000	4.90000	36.07138	36.00227	36.07137	0.06911	0.00000
52.00000	4.80000	36.00000	35.92998	35.99998	0.07002	0.00000
53.00000	4.70000	35.91063	35.83971	35.91061	0.07092	0.00000
54.00000	4.60000	35.80322	35.73135	35.80321	0.07187	0.00000
55.00000	4.50000	35.67764	35.60481	35.67763	0.07281	0.00000

Table 2. (continue)

X	INITIAL GUESS	ANALYTIC SOLUTION	SRT1 SOLUTION	SRT2 SOLUTION		
(KM)	Z ₀ (KM)	Z _A (KM)	Z ₁ (KM)	Z ₂ (KM)	Z _A - Z ₁ (KM)	Z _A - Z ₂ (KM)
56.00000	4.40000	35.53377	35.45995	35.53375	0.07382	0.00002
57.00000	4.30000	35.37146	35.29665	35.37144	0.07481	0.00002
58.00000	4.20000	35.19057	35.11473	35.19057	0.07584	0.00000
59.00000	4.10000	34.99091	34.91402	34.99088	0.07639	0.00003
60.00000	4.00000	34.77225	34.69431	34.77223	0.07794	0.00002
61.00000	3.90000	34.53438	34.45535	34.53436	0.07903	0.00002
62.00000	3.80000	34.27705	34.19691	34.27704	0.08014	0.00002
63.00000	3.70000	34.00000	33.91872	33.99997	0.08128	0.00003
64.00000	3.60000	33.70288	33.62042	33.70287	0.08246	0.00002
65.00000	3.50000	33.38539	33.30173	33.38536	0.08366	0.00003
66.00000	3.40000	33.04713	32.96225	33.04712	0.08498	0.00002
67.00000	3.30000	32.68776	32.60159	32.68773	0.08617	0.00003
68.00000	3.20000	32.30678	32.21930	32.30675	0.08748	0.00003
69.00000	3.10000	31.90375	31.81491	31.90372	0.08884	0.00003
70.00000	3.00000	31.47815	31.38788	31.47810	0.09027	0.00005
71.00000	2.90000	31.02939	30.93767	31.02936	0.09172	0.00003
72.00000	2.80000	30.55688	30.46364	30.55685	0.09325	0.00003
73.00000	2.70000	30.05995	29.95512	30.05992	0.09483	0.00003
74.00000	2.60000	29.53786	29.44138	29.53783	0.09648	0.00003
75.00000	2.50000	28.98979	28.89160	28.98975	0.09819	0.00005
76.00000	2.40000	28.41487	28.31490	28.41483	0.09998	0.00005
77.00000	2.30000	27.81212	27.71030	27.81207	0.10182	0.00005
78.00000	2.20000	27.18050	27.07674	27.18045	0.10376	0.00005
79.00000	2.10000	26.51880	26.41301	26.51875	0.10579	0.00005
80.00000	2.00000	25.82574	25.71782	25.82570	0.10793	0.00005
81.00000	1.90000	25.09988	24.98969	25.09982	0.11020	0.00006
82.00000	1.80000	24.33958	24.22701	24.33952	0.11258	0.00006
83.00000	1.70000	23.54308	23.42795	23.54301	0.11513	0.00006
84.00000	1.60000	22.70831	22.59048	22.70824	0.11783	0.00008
85.00000	1.50000	21.83249	21.71231	21.83242	0.12068	0.00008
86.00000	1.40000	20.91454	20.79076	20.91446	0.12378	0.00008
87.00000	1.30000	19.94997	19.82289	19.94987	0.12709	0.00009
88.00000	1.20000	18.93584	18.80522	18.93575	0.13062	0.00009
89.00000	1.10000	17.86818	17.73376	17.86809	0.13441	0.00009
90.00000	1.00000	16.74234	16.60382	16.74223	0.13852	0.00011
91.00000	0.90000	15.55276	15.40990	15.55266	0.14287	0.00010
92.00000	0.80000	14.29285	14.14533	14.29273	0.14752	0.00012
93.00000	0.70000	12.95451	12.80222	12.95438	0.15229	0.00013
94.00000	0.60000	11.52776	11.37091	11.52762	0.15685	0.00014
95.00000	0.50000	10.00000	9.83950	9.99984	0.16050	0.00016
96.00000	0.40000	8.35489	8.19316	8.35471	0.16173	0.00018
97.00000	0.30000	6.57065	6.41328	6.57045	0.15737	0.00020
98.00000	0.20000	4.61707	4.47622	4.61679	0.14085	0.00028
99.00000	0.10000	2.44994	2.35161	2.44977	0.09833	0.00017
100.00000	0.00000	0.00000	0.00000	0.00000	0.00000	0.00000

X	Analytic Solution Z(A)	ZH1 Solution Z	Z(A) - Z
(km)	(km)	(km)	(km)
0.00000	10.00000	10.00000	0.00000
1.00000	11.52777	11.52776	0.00001
2.00000	12.95451	12.95451	0.00000
3.00000	14.29286	14.29285	0.00001
4.00000	15.55278	15.55277	0.00001
5.00000	16.74234	16.74234	0.00000
6.00000	17.86819	17.86818	0.00002
7.00000	18.93585	18.93584	0.00002
8.00000	19.94997	19.94995	0.00002
9.00000	20.91455	20.91454	0.00002
10.00000	21.83301	21.83298	0.00003
11.00000	22.70831	22.70830	0.00002
12.00000	23.54308	23.54306	0.00002
13.00000	24.33960	24.33958	0.00002
14.00000	25.09988	25.09987	0.00002
15.00000	25.82576	25.82573	0.00003
16.00000	26.51881	26.51880	0.00002
17.00000	27.18051	27.18048	0.00003
18.00000	27.81213	27.81210	0.00003
19.00000	28.41487	28.41486	0.00002
20.00000	28.98979	28.98978	0.00002
21.00000	29.53787	29.53784	0.00003
22.00000	30.05997	30.05994	0.00003
23.00000	30.55690	30.55687	0.00003
24.00000	31.02940	31.02937	0.00003
25.00000	31.47815	31.47812	0.00003
26.00000	31.90376	31.90373	0.00003
27.00000	32.30679	32.30676	0.00003
28.00000	32.68776	32.68773	0.00003
29.00000	33.04715	33.04712	0.00003
30.00000	33.38539	33.38536	0.00003
31.00000	33.70288	33.70287	0.00002
32.00000	34.00000	33.99997	0.00003
33.00000	34.27707	34.27704	0.00003
34.00000	34.53439	34.53436	0.00003
35.00000	34.77226	34.77223	0.00003
36.00000	34.99091	34.99088	0.00003
37.00000	35.19058	35.19055	0.00003
38.00000	35.37148	35.37144	0.00003
39.00000	35.53377	35.53374	0.00003
40.00000	35.67764	35.67761	0.00003
41.00000	35.80322	35.80319	0.00003
42.00000	35.91064	35.91061	0.00003
43.00000	36.00000	35.99997	0.00003
44.00000	36.07138	36.07135	0.00003
45.00000	36.12486	36.12483	0.00003
46.00000	36.16049	36.16045	0.00005
47.00000	36.17828	36.17825	0.00003
48.00000	36.17828	36.17825	0.00003

49.00000	36.16049	36.16045	0.00005
50.00000	36.12486	36.12483	0.00003
51.00000	36.07138	36.07135	0.00003
52.00000	36.00000	35.99997	0.00003
53.00000	35.91064	35.91061	0.00003
54.00000	35.80322	35.80319	0.00003
55.00000	35.67764	35.67761	0.00003
56.00000	35.53377	35.53374	0.00003
57.00000	35.37148	35.37144	0.00003
58.00000	35.19058	35.19055	0.00003
59.00000	34.99091	34.99088	0.00003
60.00000	34.77226	34.77223	0.00003
61.00000	34.53439	34.53436	0.00003
62.00000	34.27707	34.27704	0.00003
63.00000	34.00000	33.99997	0.00003
64.00000	33.70288	33.70285	0.00003
65.00000	33.38539	33.38536	0.00003
66.00000	33.04715	33.04712	0.00003
67.00000	32.68776	32.68773	0.00003
68.00000	32.30679	32.30676	0.00003
69.00000	31.90376	31.90373	0.00003
70.00000	31.47815	31.47812	0.00003
71.00000	31.02940	31.02937	0.00003
72.00000	30.55690	30.55687	0.00003
73.00000	30.05997	30.05994	0.00003
74.00000	29.53787	29.53784	0.00003
75.00000	28.98979	28.98976	0.00003
76.00000	28.41487	28.41484	0.00003
77.00000	27.81213	27.81210	0.00003
78.00000	27.18051	27.18048	0.00003
79.00000	26.51881	26.51878	0.00003
80.00000	25.82576	25.82573	0.00003
81.00000	25.09988	25.09987	0.00002
82.00000	24.33960	24.33957	0.00003
83.00000	23.54308	23.54306	0.00002
84.00000	22.70831	22.70828	0.00003
85.00000	21.83301	21.83298	0.00003
86.00000	20.91455	20.91452	0.00003
87.00000	19.94997	19.94995	0.00002
88.00000	18.93585	18.93582	0.00003
89.00000	17.86819	17.86816	0.00003
90.00000	16.74234	16.74232	0.00002
91.00000	15.55278	15.55276	0.00002
92.00000	14.29286	14.29284	0.00002
93.00000	12.95451	12.95450	0.00001
94.00000	11.52777	11.52775	0.00002
95.00000	10.00000	9.99999	0.00001
96.00000	8.35489	8.35488	0.00001
97.00000	6.57066	6.57065	0.00001
98.00000	4.61707	4.61706	0.00001
99.00000	2.44995	2.44994	0.00001
100.00000	0.00000	-0.00000	0.00000

8.2 Test Case No.2

In this case we assume $V(x,z) = 4 + 0.01 x + 0.2 z$ (km/sec). The two end points are the same as in case No.1, i.e. (0, 10) and (100, 0) (km). The analytical solution is a circular ray path with its center at (47.263683, -22.363159) (km). In Table 4, the computed solution from the programs SRT1 and SRT2 (Lee's) are shown. In Table 5 the computed solution from program ZH1 together with the analytic solution are shown. The maximum deviations of the SRT1 and SRT2 solutions from the analytic solution are about 0.14309 km and 0.00014 km respectively; whereas the maximum deviation of the ZH1 solution from the analytic one is about 0.00005 km. In this example we once again see that the accuracy of program ZH1 is better than that of programs SRT1 and SRT2.

Table 4. Comparison of SRT1 and SRT2 Solutions for Case No.2

	INITIAL GUESS	SRT1 SOLUTION	SRT2 SOLUTION	
X	Z ₀	Z ₁	Z ₂	Z ₁ - Z ₂
(KM)	(KM)	(KM)	(KM)	(KM)
0.00000	10.00000	10.00000	10.00000	0.00000
1.00000	9.90000	11.40590	11.41407	-0.00817
2.00000	9.80000	12.72872	12.74230	-0.01408
3.00000	9.70000	13.97702	13.99555	-0.01853
4.00000	9.60000	15.15790	15.17992	-0.02202
5.00000	9.50000	16.27733	16.30220	-0.02487
6.00000	9.40000	17.34042	17.36765	-0.02722
7.00000	9.30000	18.35146	18.38074	-0.02928
8.00000	9.20000	19.31421	19.34526	-0.03105
9.00000	9.10000	20.23187	20.26453	-0.03265
10.00000	9.00000	21.10732	21.14140	-0.03409
11.00000	8.90000	21.94301	21.97839	-0.03539
12.00000	8.80000	22.74113	22.77774	-0.03661
13.00000	8.70000	23.50365	23.54140	-0.03775
14.00000	8.60000	24.22230	24.27112	-0.03882
15.00000	8.50000	24.92863	24.96848	-0.03984
16.00000	8.40000	25.59407	25.63487	-0.04080
17.00000	8.30000	26.22986	26.27158	-0.04172
18.00000	8.20000	26.83716	26.87976	-0.04260
19.00000	8.10000	27.41701	27.46045	-0.04344
20.00000	8.00000	27.97035	28.01460	-0.04425
21.00000	7.90000	28.49805	28.54311	-0.04506
22.00000	7.80000	29.00089	29.04671	-0.04582
23.00000	7.70000	29.47958	29.52615	-0.04657
24.00000	7.60000	29.93480	29.98213	-0.04733
25.00000	7.50000	30.36716	30.41521	-0.04805
26.00000	7.40000	30.77719	30.82597	-0.04878
27.00000	7.30000	31.16542	31.21492	-0.04950
28.00000	7.20000	31.53232	31.58253	-0.05022
29.00000	7.10000	31.87833	31.92924	-0.05092
30.00000	7.00000	32.20383	32.25545	-0.05162
31.00000	6.90000	32.50919	32.56152	-0.05234
32.00000	6.80000	32.79475	32.84778	-0.05302
33.00000	6.70000	33.06081	33.11453	-0.05373
34.00000	6.60000	33.30765	33.36208	-0.05443
35.00000	6.50000	33.53554	33.59065	-0.05511
36.00000	6.40000	33.74469	33.80051	-0.05582
37.00000	6.30000	33.93532	33.99182	-0.05650
38.00000	6.20000	34.10760	34.16481	-0.05721
39.00000	6.10000	34.26172	34.31963	-0.05791
40.00000	6.00000	34.39783	34.45644	-0.05861
41.00000	5.90000	34.51604	34.57533	-0.05930
42.00000	5.80000	34.61647	34.67648	-0.06001
43.00000	5.70000	34.69920	34.75993	-0.06073
44.00000	5.60000	34.76433	34.82578	-0.06145
45.00000	5.50000	34.81189	34.87408	-0.06219
46.00000	5.40000	34.84195	34.90489	-0.06294
47.00000	5.30000	34.85454	34.91823	-0.06369
48.00000	5.20000	34.84964	34.91409	-0.06445
49.00000	5.10000	34.82726	34.89252	-0.06526
50.00000	5.00000	34.78738	34.85344	-0.06606
51.00000	4.90000	34.73000	34.79684	-0.06685
52.00000	4.80000	34.65507	34.72269	-0.06767
53.00000	4.70000	34.56238	34.63087	-0.06850
54.00000	4.60000	34.45201	34.52136	-0.06935
55.00000	4.50000	34.32381	34.39400	-0.07019

Table 4. (continue)

X	INITIAL GUESS	SRT1 SOLUTION	SRT2 SOLUTION	
(KM)	Z ₀ (KM)	Z ₁ (KM)	Z ₂ (KM)	Z ₁ - Z ₂ (KM)
56.00000	4.40000	34.17763	34.24870	-0.07108
57.00000	4.30000	34.01334	34.08531	-0.07198
58.00000	4.20000	33.83081	33.90369	-0.07288
59.00000	4.10000	33.62982	33.70363	-0.07381
60.00000	4.00000	33.41020	33.48495	-0.07475
61.00000	3.90000	33.17172	33.24745	-0.07573
62.00000	3.80000	32.91414	32.99084	-0.07671
63.00000	3.70000	32.63719	32.71490	-0.07771
64.00000	3.60000	32.34056	32.41933	-0.07877
65.00000	3.50000	32.02396	32.10379	-0.07983
66.00000	3.40000	31.68701	31.76794	-0.08093
67.00000	3.30000	31.32935	31.41141	-0.08206
68.00000	3.20000	30.95055	31.03375	-0.08321
69.00000	3.10000	30.55016	30.63457	-0.08441
70.00000	3.00000	30.12767	30.21330	-0.08563
71.00000	2.90000	29.68257	29.76949	-0.08691
72.00000	2.80000	29.21426	29.30249	-0.08823
73.00000	2.70000	28.72211	28.81171	-0.08960
74.00000	2.60000	28.20541	28.29643	-0.09102
75.00000	2.50000	27.66344	27.75592	-0.09248
76.00000	2.40000	27.09531	27.18933	-0.09402
77.00000	2.30000	26.50012	26.59576	-0.09564
78.00000	2.20000	25.87689	25.97423	-0.09734
79.00000	2.10000	25.22453	25.32364	-0.09911
80.00000	2.00000	24.54181	24.64278	-0.10097
81.00000	1.90000	23.82738	23.93031	-0.10294
82.00000	1.80000	23.07974	23.18475	-0.10501
83.00000	1.70000	22.29726	22.40446	-0.10721
84.00000	1.60000	21.47803	21.58759	-0.10956
85.00000	1.50000	20.62000	20.73203	-0.11203
86.00000	1.40000	19.72076	19.83543	-0.11467
87.00000	1.30000	18.77765	18.89516	-0.11751
88.00000	1.20000	17.78758	17.90813	-0.12054
89.00000	1.10000	16.74699	16.87080	-0.12381
90.00000	1.00000	15.65183	15.77909	-0.12726
91.00000	0.90000	14.49728	14.62817	-0.13089
92.00000	0.80000	13.27774	13.41233	-0.13459
93.00000	0.70000	11.98652	12.12471	-0.13819
94.00000	0.60000	10.61568	10.75693	-0.14125
95.00000	0.50000	9.15562	9.29861	-0.14299
96.00000	0.40000	7.59470	7.73658	-0.14188
97.00000	0.30000	5.91865	6.05377	-0.13512
98.00000	0.20000	4.10985	4.22722	-0.11737
99.00000	0.10000	2.14627	2.22503	-0.07876
100.00000	0.00000	0.00000	0.00000	0.00000

X	Analytic Solution Z(A)	ZH1 Solution Z	Z(A) - Z
(km)	(km)	(km)	(km)
0.00000	10.00000	10.00000	0.00000
1.00000	11.41407	11.41407	0.00000
2.00000	12.74281	12.74280	0.00000
3.00000	13.99555	13.99555	0.00000
4.00000	15.17993	15.17992	0.00001
5.00000	16.30220	16.30220	0.00000
6.00000	17.36766	17.36765	0.00002
7.00000	18.38074	18.38072	0.00002
8.00000	19.34526	19.34525	0.00002
9.00000	20.26453	20.26451	0.00002
10.00000	21.14140	21.14139	0.00002
11.00000	21.97841	21.97839	0.00002
12.00000	22.77776	22.77773	0.00003
13.00000	23.54141	23.54138	0.00003
14.00000	24.27113	24.27110	0.00003
15.00000	24.96848	24.96846	0.00002
16.00000	25.63487	25.63486	0.00002
17.00000	26.27159	26.27156	0.00003
18.00000	26.87976	26.87975	0.00002
19.00000	27.46046	27.46043	0.00003
20.00000	28.01462	28.01459	0.00003
21.00000	28.54311	28.54308	0.00003
22.00000	29.04672	29.04669	0.00003
23.00000	29.52617	29.52614	0.00003
24.00000	29.98213	29.98210	0.00003
25.00000	30.41522	30.41519	0.00003
26.00000	30.82597	30.82596	0.00002
27.00000	31.21494	31.21490	0.00003
28.00000	31.58255	31.58252	0.00003
29.00000	31.92926	31.92923	0.00003
30.00000	32.25546	32.25543	0.00003
31.00000	32.56152	32.56149	0.00003
32.00000	32.84779	32.84776	0.00003
33.00000	33.11455	33.11452	0.00003
34.00000	33.36209	33.36206	0.00003
35.00000	33.59067	33.59064	0.00003
36.00000	33.80051	33.80048	0.00003
37.00000	33.99184	33.99181	0.00003
38.00000	34.16483	34.16479	0.00003
39.00000	34.31964	34.31961	0.00003
40.00000	34.45645	34.45642	0.00003
41.00000	34.57536	34.57533	0.00003
42.00000	34.67650	34.67647	0.00003
43.00000	34.75995	34.75992	0.00003
44.00000	34.82581	34.82576	0.00005
45.00000	34.87411	34.87407	0.00005
46.00000	34.90492	34.90488	0.00005
47.00000	34.91824	34.91821	0.00003
48.00000	34.91412	34.91408	0.00005

49.00000	34.89253	34.89250	0.00003
50.00000	34.85345	34.85342	0.00003
51.00000	34.79687	34.79683	0.00005
52.00000	34.72272	34.72267	0.00005
53.00000	34.63091	34.63087	0.00003
54.00000	34.52138	34.52135	0.00003
55.00000	34.39403	34.39400	0.00003
56.00000	34.24873	34.24869	0.00005
57.00000	34.08534	34.08530	0.00005
58.00000	33.90372	33.90367	0.00005
59.00000	33.70366	33.70361	0.00005
60.00000	33.48499	33.48495	0.00003
61.00000	33.24747	33.24744	0.00003
62.00000	32.99088	32.99084	0.00003
63.00000	32.71494	32.71490	0.00003
64.00000	32.41936	32.41933	0.00003
65.00000	32.10382	32.10379	0.00003
66.00000	31.76797	31.76793	0.00005
67.00000	31.41144	31.41139	0.00005
68.00000	31.03378	31.03375	0.00003
69.00000	30.63458	30.63455	0.00003
70.00000	30.21335	30.21332	0.00003
71.00000	29.76952	29.76949	0.00003
72.00000	29.30254	29.30249	0.00005
73.00000	28.81174	28.81171	0.00003
74.00000	28.29648	28.29643	0.00005
75.00000	27.75595	27.75592	0.00003
76.00000	27.18936	27.18933	0.00003
77.00000	26.59581	26.59578	0.00003
78.00000	25.97427	25.97424	0.00003
79.00000	25.32368	25.32365	0.00003
80.00000	24.64282	24.64279	0.00003
81.00000	23.93036	23.93033	0.00003
82.00000	23.18480	23.18478	0.00002
83.00000	22.40451	22.40448	0.00003
84.00000	21.58763	21.58760	0.00003
85.00000	20.73207	20.73206	0.00002
86.00000	19.83549	19.83548	0.00002
87.00000	18.89522	18.89520	0.00002
88.00000	17.90819	17.90817	0.00002
89.00000	16.87088	16.87085	0.00003
90.00000	15.77917	15.77914	0.00003
91.00000	14.62825	14.62823	0.00002
92.00000	13.41241	13.41239	0.00002
93.00000	12.12479	12.12478	0.00002
94.00000	10.75702	10.75700	0.00002
95.00000	9.29871	9.29869	0.00002
96.00000	7.73669	7.73668	0.00001
97.00000	6.05389	6.05388	0.00001
98.00000	4.22736	4.22736	0.00001
99.00000	2.22512	2.22511	0.00001
100.00000	0.00000	-0.00000	0.00000

8.3 Summary of Test Case Results

In Table 6 the main results of our test cases are summarized using some solutions of Lee's. We shall show the maximal deviation from the analytic solutions or from each other and the total travel time given by each program. It is clear from Table 6 that all programs SRT1, SRT2 and ZH1 will compute the travel time correctly. Finally, by comparison with analytic solutions, we note that the accuracy of program ZH1 is better than those of programs SRT1 and SRT2. In other words, Pereyra's method is more accurate than Wesson's and Lentini's methods.

Table 6. Comparison of Analytic, SRT1, SRT2 and ZH1 Solutions

Case No.	Velocity Model (km/sec)	Solution	Travel Time (sec)	Maximal Deviation (km)
1.	4 + 0.2 Z	Analytic	14.6640	
	A(0, 10)	SRT1	14.6625	0.16 (SRT1 - Anal.)
	B(100, 0)	SRT2	14.6624	0.00028 (SRT2 - Anal.)
		ZH1	14.6635	0.00003 (ZH1 - Anal.)
2.	4 + 0.01 X + 0.2 Z	Analytic	13.666	
	A(0, 10)	SRT1	13.666	0.14309 (SRT1 - Anal.)
	B(100, 0)	SRT2	13.666	0.00014 (SRT2 - Anal.)
		ZH1	13.666	0.00005 (ZH1 - Anal.)

8.4 The Program ZH1 and 5 Computed Examples of Ray Tracing Problems in 2-D Media

In 8.3 section we have compared the results of program ZH1 with Lee's programs SRT1 and SRT2. In order to show the computed efficiency of ZH1 we have used it on MTS computer to compute results using 5 different geological models. The CPU times needed are shown in Table 7. In the Tables 8 and 12 the final output solutions including the velocity model, the total mesh number N used, and the expected accuracy TOL, are all given. We shall see from Table 7 that the computation efficiency of ZH1 is high reasonably and program ZH1 may be used as a useful subroutine in any large program of ray approximation which needs to compute the ray path and the travel time.

The program ZH1 is listed in Appendix I and is written in ALGOLW language and is valid on the MTS computer at the University of Alberta. A user will be asked only to input the velocity model and the desired accuracy.

Table 7. Main Computed Solutions Using ZH1 Program

No.	Velocity Model (km/sec)	Total Travel Time (sec)	Total Ray Length (km)	CPU Time on MTS (sec)
1.	4 + 0.01 X + 0.2 Z A(0, 10); B(60, 20)	7.687	65.828	0.132
2.	4 + 0.01 X + 0.2 Z A(0, 0); B(50, 0)	10.018	56.830	0.134
3.	5 + 0.05 X + 0.1 Z A(20, 20); B(60, 0)	5.503	45.411	0.099
4.	4 + 0.01 X + 0.05 Z A(30, 15); B(60, 5)	6.362	31.758	0.066
5.	6 + 0.05 X + 0.1 Z A(5, 3); B(65, 10)	7.536	61.580	0.104

Table 8 . Computed Solution of ZH1 for Model 1 .

VELOCITY MODEL:

$$V = Z_0 + Z_1 * X + Z_2 * Z$$

$$Z_0 = 4.000000 \quad Z_1 = 0.0100000 \quad Z_2 = 0.2000000$$

$$X(0) = 0 \quad Z(0) = 10.00000$$

$$X(1) = 60.00000 \quad Z(1) = 20.00000$$

$$N = 21$$

$$TOL = 1.00 \times 10^{-3} \quad (7.67 \times 10^{-5})$$

N= 1	X=	0.000	Z=	10.000	Q=	48.646	T=	0.000	S=	0.000
N= 2	X=	3.000	Z=	13.114	Q=	43.496	T=	0.684	S=	4.324
N= 3	X=	6.000	Z=	15.734	Q=	38.755	T=	1.258	S=	8.307
N= 4	X=	9.000	Z=	17.956	Q=	34.310	T=	1.760	S=	12.040
N= 5	X=	12.000	Z=	19.845	Q=	30.089	T=	2.209	S=	15.586
N= 6	X=	15.000	Z=	21.445	Q=	26.042	T=	2.621	S=	18.985
N= 7	X=	18.000	Z=	22.786	Q=	22.130	T=	3.003	S=	22.272
N= 8	X=	21.000	Z=	23.891	Q=	18.324	T=	3.364	S=	25.469
N= 9	X=	24.000	Z=	24.778	Q=	14.600	T=	3.708	S=	28.597
N= 10	X=	27.000	Z=	25.458	Q=	10.938	T=	4.040	S=	31.673
N= 11	X=	30.000	Z=	25.940	Q=	7.321	T=	4.362	S=	34.712
N= 12	X=	33.000	Z=	26.230	Q=	3.733	T=	4.678	S=	37.726
N= 13	X=	36.000	Z=	26.332	Q=	0.160	T=	4.991	S=	40.727
N= 14	X=	39.000	Z=	26.247	Q=	-3.413	T=	5.303	S=	43.728
N= 15	X=	42.000	Z=	25.974	Q=	-6.999	T=	5.615	S=	46.741
N= 16	X=	45.000	Z=	25.509	Q=	-10.613	T=	5.932	S=	49.777
N= 17	X=	48.000	Z=	24.847	Q=	-14.270	T=	6.256	S=	52.849
N= 18	X=	51.000	Z=	23.980	Q=	-17.987	T=	6.589	S=	55.972
N= 19	X=	54.000	Z=	22.894	Q=	-21.785	T=	6.935	S=	59.162
N= 20	X=	57.000	Z=	21.575	Q=	-25.686	T=	7.299	S=	62.439
N= 21	X=	60.000	Z=	20.000	Q=	-29.720	T=	7.687	S=	65.828

Table 9. Computed Solution of ZH1 for Model 2.

VELOCITY MODEL:

$$V = Z_0 + Z_1 * X + Z_2 * Z$$

$$Z_0 = 4.000000 \quad Z_1 = 0.0100000 \quad Z_2 = 0.2000000$$

$$X(0) = 0 \quad Z(0) = 0$$

$$X(1) = 50.00000 \quad Z(1) = 0$$

$$N = 21$$

$$TOL = 1.00 \times 10^{-3} \quad (4.86 \times 10^{-4})$$

N= 1	X=	0.000	Z=	0.000	Q=	49.640	T=	0.000	S=	0.000
N= 2	X=	2.500	Z=	2.631	Q=	43.294	T=	0.849	S=	3.630
N= 3	X=	5.000	Z=	4.761	Q=	37.559	T=	1.536	S=	6.914
N= 4	X=	7.500	Z=	6.505	Q=	32.233	T=	2.124	S=	9.962
N= 5	X=	10.000	Z=	7.932	Q=	27.204	T=	2.644	S=	12.840
N= 6	X=	12.500	Z=	9.087	Q=	22.394	T=	3.118	S=	15.594
N= 7	X=	15.000	Z=	10.000	Q=	17.745	T=	3.558	S=	18.256
N= 8	X=	17.500	Z=	10.693	Q=	13.213	T=	3.974	S=	20.850
N= 9	X=	20.000	Z=	11.178	Q=	8.765	T=	4.374	S=	23.397
N= 10	X=	22.500	Z=	11.466	Q=	4.370	T=	4.762	S=	25.913
N= 11	X=	25.000	Z=	11.561	Q=	0.000	T=	5.145	S=	28.415
N= 12	X=	27.500	Z=	11.466	Q=	-4.370	T=	5.526	S=	30.917
N= 13	X=	30.000	Z=	11.178	Q=	-8.765	T=	5.910	S=	33.433
N= 14	X=	32.500	Z=	10.693	Q=	-13.213	T=	6.302	S=	35.980
N= 15	X=	35.000	Z=	10.000	Q=	-17.744	T=	6.707	S=	38.574
N= 16	X=	37.500	Z=	9.087	Q=	-22.394	T=	7.131	S=	41.236
N= 17	X=	40.000	Z=	7.932	Q=	-27.204	T=	7.584	S=	43.990
N= 18	X=	42.500	Z=	6.505	Q=	-32.233	T=	8.075	S=	46.868
N= 19	X=	45.000	Z=	4.761	Q=	-37.559	T=	8.623	S=	49.917
N= 20	X=	47.500	Z=	2.631	Q=	-43.294	T=	9.254	S=	53.201
N= 21	X=	50.000	Z=	-0.000	Q=	-49.640	T=	10.018	S=	56.830

Table 10. Computed Solution of ZH1 for Model 3.

VELOCITY MODEL:

$$V = Z_0 + Z_1 * X + Z_2 * Z$$

$$Z_0 = 5.000000 \quad Z_1 = 0.0500000 \quad Z_2 = 0.1000000$$

$$X(0) = 20.00000 \quad Z(0) = 20.00000$$

$$X(1) = 60.00000 \quad Z(1) = 0$$

$$N = 21$$

$$TOL = 1.00 \times 10^{-3} \quad (3.87 \times 10^{-5})$$

N= 1	X=	20.000	Z=	20.000	Q=	-9.211	T=	0.000	S=	0.000
N= 2	X=	22.000	Z=	19.648	Q=	-10.763	T=	0.253	S=	2.031
N= 3	X=	24.000	Z=	19.239	Q=	-12.323	T=	0.505	S=	4.072
N= 4	X=	26.000	Z=	18.774	Q=	-13.893	T=	0.757	S=	6.126
N= 5	X=	28.000	Z=	18.250	Q=	-15.473	T=	1.009	S=	8.193
N= 6	X=	30.000	Z=	17.666	Q=	-17.066	T=	1.262	S=	10.277
N= 7	X=	32.000	Z=	17.021	Q=	-18.672	T=	1.515	S=	12.378
N= 8	X=	34.000	Z=	16.314	Q=	-20.293	T=	1.770	S=	14.499
N= 9	X=	36.000	Z=	15.541	Q=	-21.932	T=	2.027	S=	16.643
N= 10	X=	38.000	Z=	14.702	Q=	-23.589	T=	2.287	S=	18.812
N= 11	X=	40.000	Z=	13.794	Q=	-25.268	T=	2.549	S=	21.009
N= 12	X=	42.000	Z=	12.813	Q=	-26.971	T=	2.815	S=	23.236
N= 13	X=	44.000	Z=	11.757	Q=	-28.699	T=	3.085	S=	25.498
N= 14	X=	46.000	Z=	10.622	Q=	-30.457	T=	3.360	S=	27.798
N= 15	X=	48.000	Z=	9.403	Q=	-32.247	T=	3.640	S=	30.140
N= 16	X=	50.000	Z=	8.097	Q=	-34.073	T=	3.927	S=	32.529
N= 17	X=	52.000	Z=	6.696	Q=	-35.939	T=	4.222	S=	34.970
N= 18	X=	54.000	Z=	5.195	Q=	-37.851	T=	4.525	S=	37.471
N= 19	X=	56.000	Z=	3.585	Q=	-39.813	T=	4.838	S=	40.039
N= 20	X=	58.000	Z=	1.857	Q=	-41.833	T=	5.164	S=	42.681
N= 21	X=	60.000	Z=	0.000	Q=	-43.919	T=	5.503	S=	45.411

Table 11. Computed Solution of ZH1 for Model 4.

VELOCITY MODEL:

$V = Z_0 + Z_1 * X + Z_2 * Z$
 $Z_0 = 4.000000 \quad Z_1 = 0.0100000 \quad Z_2 = 0.0500000$

$X(0) = 30.00000 \quad Z(0) = 15.00000$
 $X(1) = 60.00000 \quad Z(1) = 5.000000$

$N = 21$
 $TOL = 1.00 \times 10^{-3} \quad (2.19 \times 10^{-4})$

N=	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
N=	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21
X=	30.000	31.500	33.000	34.500	36.000	37.500	39.000	40.500	42.000	43.500	45.000	46.500	48.000	49.500	51.000	52.500	54.000	55.500	57.000	58.500	60.000
Z=	15.000	14.744	14.464	14.160	13.831	13.479	13.101	12.699	12.271	11.817	11.337	10.831	10.298	9.738	9.149	8.532	7.887	7.211	6.505	5.769	5.000
Q=	-9.252	-10.132	-11.014	-11.899	-12.787	-13.678	-14.573	-15.470	-16.372	-17.278	-18.189	-19.104	-20.025	-20.951	-21.882	-22.820	-23.765	-24.716	-25.674	-26.641	-27.615
T=	0.000	0.301	0.603	0.906	1.210	1.515	1.822	2.130	2.439	2.751	3.065	3.380	3.699	4.020	4.344	4.671	5.001	5.335	5.673	6.015	6.362
S=	0.000	1.522	3.048	4.578	6.114	7.655	9.201	10.754	12.314	13.881	15.456	17.039	18.631	20.232	21.844	23.466	25.099	26.744	28.402	30.073	31.758

Table 12. Computed Solution of ZH1 for Model 5.

VELOCITY MODEL:

V=Z0+Z1*X+Z2*Z
Z0= 5.599999 Z1= 0.050000 Z2= 0.100000

X(0)= 5.000000 Z(0)= 3.000000
X(1)= 65.00000 Z(1)= 10.00000

N= 21
TOL= 1.00E-03 (1.51E-05)

N= 1	X= 5.000	Z= 3.000	Q= 26.104	T=	0.000	S= 0.000
N= 2	X= 8.000	Z= 4.403	Q= 24.012	T=	0.526	S= 3.312
N= 3	X= 11.000	Z= 5.675	Q= 21.953	T=	1.021	S= 6.570
N= 4	X= 14.000	Z= 6.823	Q= 19.924	T=	1.490	S= 9.782
N= 5	X= 17.000	Z= 7.851	Q= 17.921	T=	1.936	S= 12.954
N= 6	X= 20.000	Z= 8.765	Q= 15.940	T=	2.363	S= 16.090
N= 7	X= 23.000	Z= 9.566	Q= 13.979	T=	2.772	S= 19.195
N= 8	X= 26.000	Z= 10.259	Q= 12.034	T=	3.166	S= 22.274
N= 9	X= 29.000	Z= 10.846	Q= 10.103	T=	3.546	S= 25.331
N= 10	X= 32.000	Z= 11.329	Q= 8.183	T=	3.915	S= 28.369
N= 11	X= 35.000	Z= 11.709	Q= 6.273	T=	4.274	S= 31.393
N= 12	X= 38.000	Z= 11.989	Q= 4.370	T=	4.624	S= 34.406
N= 13	X= 41.000	Z= 12.168	Q= 2.471	T=	4.966	S= 37.412
N= 14	X= 44.000	Z= 12.248	Q= 0.576	T=	5.302	S= 40.413
N= 15	X= 47.000	Z= 12.228	Q= -1.319	T=	5.632	S= 43.413
N= 16	X= 50.000	Z= 12.110	Q= -3.216	T=	5.956	S= 46.415
N= 17	X= 53.000	Z= 11.891	Q= -5.116	T=	6.277	S= 49.423
N= 18	X= 56.000	Z= 11.572	Q= -7.022	T=	6.595	S= 52.440
N= 19	X= 59.000	Z= 11.152	Q= -8.935	T=	6.910	S= 55.469
N= 20	X= 62.000	Z= 10.628	Q= -10.859	T=	7.224	S= 58.515
N= 21	X= 65.000	Z= 10.000	Q= -12.795	T=	7.536	S= 61.580

8.5 A Simple Computed Example for 2-D Model With an Interface

We have programed Eq.(85) with the related boundary conditions (88) which expresses two point ray tracing problem in a 2-D model with an interface. Using this program we computed quantities for the following model.

$$V_1(x,z) = 8.5 + 0.04 x + 0.05 z \text{ (km/sec),}$$

$$V_2(x,z) = 3.0 + 0.05 x + 0.2 z \text{ (km/sec),}$$

$$A(0,60) \text{ and } B(30,0) \text{ (km),}$$

$$\text{Interface: } z = 50 \text{ (km).}$$

The final computed solutions are:

(1) The position of incident point C is:

$$(x = 7.978 \text{ km, } z = 50.000 \text{ km}).$$

(2) The incident angle is:

$$\theta_i = 36.36^\circ.$$

(3) The reflected angle is:

$$\theta_r = 44.51^\circ.$$

(4) The length of path in region (1) is:

$$S_{(1)} = 12.795 \text{ km.}$$

(5) The length of path in region (2) is:

$$S_{(2)} = 55.889 \text{ km.}$$

(6) The total length of path is:

$$S = S_{(1)} + S_{(2)} = 68.684 \text{ km.}$$

(7) The travel time in region (1) is:

$$T_{(1)} = 1.121 \text{ sec.}$$

(8) The travel time in region (2) is:

$$T_{(2)} = 6.531 \text{ sec.}$$

(9) The Total travel time is:

$$T = T_{(1)} + T_{(2)} = 7.652 \text{ sec.}$$

Using the incident angle, reflected angle, $V_1(c) = 13.399$ km/sec and $V_2(c) = 11.319$ km/sec we have:

$$\sin\theta_i/V_1(c) = 0.0527 = \sin\theta_r/V_2(c),$$

which shows us that the computed solutions satisfy Snell's law at the incident point c. This means that the solution computed by our program is correct.

9. Method for Computation of Ray Amplitudes

From Eq.(40) or (42) we can see that the most complicated part in the computation of ray amplitudes is the evaluation of the geometrical spreading, which can be simply expressed in terms of the Jacobian, J , of the transformation from Cartesian coordinates (x,y,z) into ray coordinates (s,q_2,q_3) .

In what follows, we shall be interested mainly in the evaluation of the spreading function J , which can be computed in several ways. V. Cervený et al (1979) have summarized some methods which are based on the computation of the Jacobian, J , at a number of points along the ray. The second or third derivative of the velocity is usually required in those schemes. It is obvious that the numerical computations can become rather tedious using the above methods.

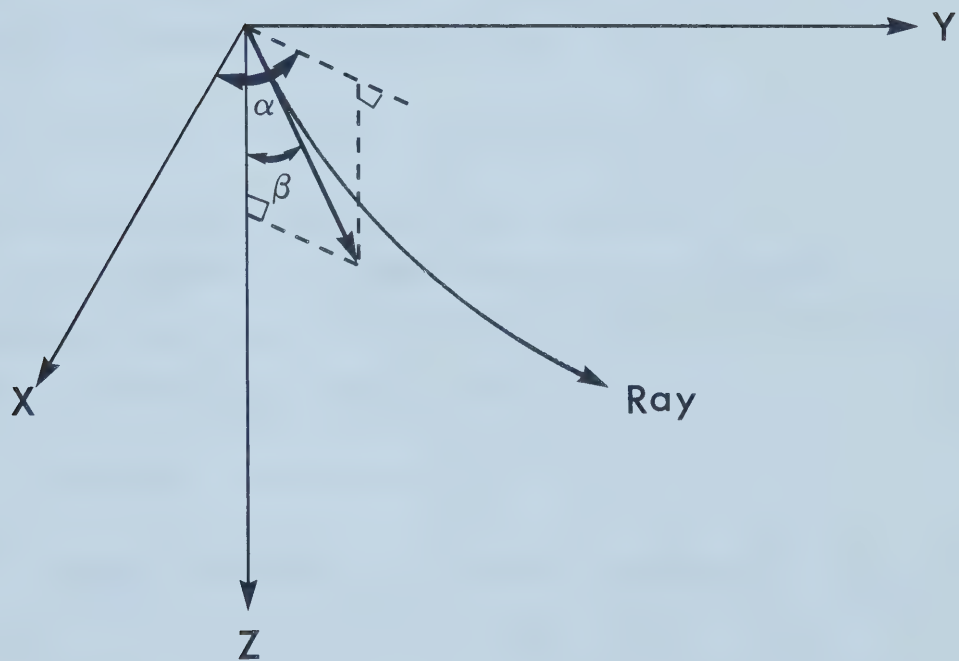
Here we shall present a new method of calculating the Jacobian, J . We are interested in solving J based on the ray tracing system Eq.(63) - (66). As shown in Fig.4, taking q_2 and q_3 as the taking off angle β and the azimuth angle α respectively, we have:

$$(95) \quad \Delta\beta = \mp \Delta(dz/ds) / (1 - (dz/ds)^2)^{1/2},$$

$$(96) \quad \Delta\alpha = [(dx/ds)\Delta(dy/ds) - (dy/ds)\Delta(dx/ds)] \\ / [(dx/ds)^2 + (dy/ds)^2],$$

where, " Δ " indicates the small increase.

FIGURE 4. Definition of the taking off angle β and the azimuth angle α .



We may introduce the following perturbations in two different fashions.

At the taking off point (source) of the ray under consideration, we increase the quantity dz/ds by the small increment $\Delta(dz/ds)$ and as a result (dx/ds) and (dy/ds) become $\lambda(dx/ds)$ and $\lambda(dy/ds)$, respectively, where, λ is determined according to Eq.(65) as:

$$(97) \quad [dz/ds + \Delta(dz/ds)]^2 + \lambda^2[(dx/ds)^2 + (dy/ds)^2] = 1.$$

This only produces a change in the angle β .

At the taking off point, we introduce the small incremental quantities $\Delta(dx/ds)$ and $\Delta(dy/ds)$ into (dx/ds) and (dy/ds) , respectively, which gives:

$$(98) \quad \begin{aligned} \Delta(dy/ds) &= \lambda\Delta(dx/ds), \\ [dx/ds + \Delta(dx/ds)]^2 + [dy/ds + \lambda\Delta(dx/ds)]^2 \\ &= (dx/ds)^2 + (dy/ds)^2. \end{aligned}$$

This only produces a change in the angle α .

Upon solving the ray tracing system Eq.(63)-(66), We also get at the same time the following values at the source: $dx(0)/ds$, $dy(0)/ds$, $dz(0)/ds$. If we change the values of $dx(0)/ds$, $dy(0)/ds$, $dz(0)/ds$ by introducing the small perturbations as mentioned above. we will get two new groups of initial values $dx(0)/ds|_i$, $dy(0)/ds|_i$ and $dz(0)/ds|_i$ ($i = 2$ or 3) and form the following initial value problems which describes the perturbed rays using the same symbols as in Eq.(63) - (66):

$$\begin{aligned}
 dY_1/ds &= Y_2, \\
 dY_2/ds &= V[u_x - GY_2], \\
 dY_3/ds &= Y_4, \\
 dY_4/ds &= V[u_y - GY_4], \\
 dY_5/ds &= Y_6, \\
 dY_6/ds &= V[u_z - GY_6],
 \end{aligned}
 \tag{99}$$

with the related initial value conditions:

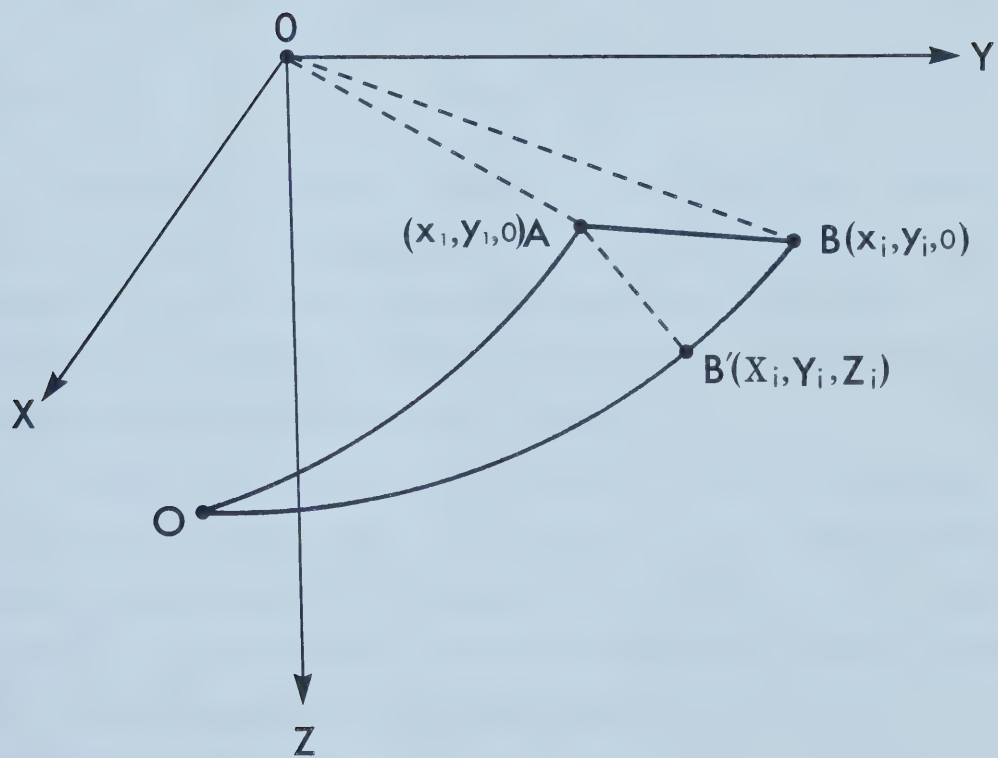
$$\begin{aligned}
 Y_1(0) &= x_0, & Y_2(0) &= dx(0)/ds|_i, \\
 Y_3(0) &= y_0, & Y_4(0) &= dy(0)/ds|_i, & (i = 2 \text{ or } 3) \\
 Y_5(0) &= z_0, & Y_6(0) &= dz(0)/ds|_i.
 \end{aligned}
 \tag{100}$$

Solving Eq.(99) and (100) we can get the coordinates of the perturbed rays at the surface $(x_i, y_i, 0)$, where, $i = 2$ or 3 .

As shown in Fig.5, the point O denotes the source and OA is the ray under consideration while OB_i ($i=2,3$) are the perturbed rays. The points A and B_i ($i=2,3$) are all placed at the surface, with $A(x_i, y_i, 0)$ being the receiver point.

It is obvious that the point $B(x_i, y_i, 0)$, where $i = 2$ or 3 , is not usually placed on the wavefront plane which passes through the receiver point A.

FIGURE 5. The ray under consideration OA and the perturbed ray OB.



In the following, using elementary geometrical reasoning we shall consider the coordinate of the point B' , $x(S)_i$, $y(S)_i$, $z(S)_i$ ($i = 2$ or 3), where the perturbed ray intersects with the wavefront plane which passes through the receiver point, A .

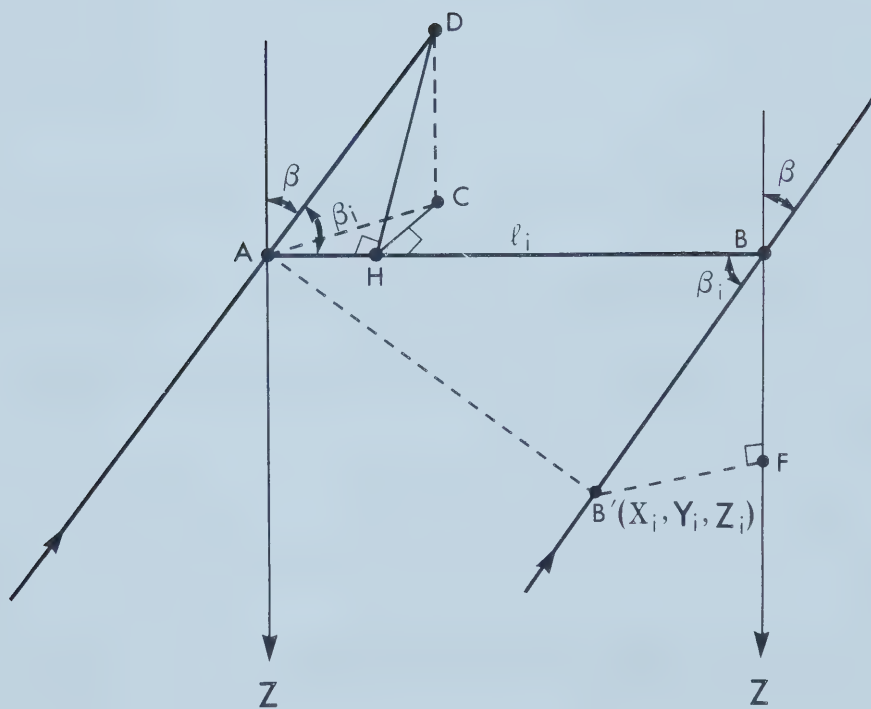
Writing $|AB| = \ell_i$ we have:

$$(101) \quad \ell_i = [(x_i - x_1)^2 + (y_i - y_1)^2]^{1/2}, \quad i = 2 \text{ or } 3.$$

As shown in Fig.6, at point A , let's define D so that $|AD| = 1$ along the tangential direction to the ray OA and assume β is the acute angle between AD and the z -axis. Making DC parallel to the z -axis and CH perpendicular to AB , we then have DH perpendicular to AB .

Assume $\beta + d\beta$ is the acute angle between the perturbed ray OB and z -coordinate. Although $d\beta$ is not usually equal to zero, we may neglect it because its contribution is a small quantity of second order. So we may consider that AB , AD and BB' are all placed on the same plane and AB' is perpendicular to BB' .

FIGURE 6. Explanation of various symbols for the computation of the quantity of BB' .



As shown in Fig.7, letting $\alpha(AB)$ be the angle between AB and x-coordinate, we have:

$$(102) \quad \tan \alpha(AB) = (y_i - y_1) / (x_i - x_1), \quad i = 2 \text{ or } 3.$$

By definition of α we may write the angle between CA and x-coordinate as α_1 . If we denote the angle between CA and AB as α_i , we have:

$$(103) \quad \alpha_i = \alpha_1 - \alpha(AB),$$

$$\tan \alpha_i = [\tan \alpha_1 - \tan \alpha(AB)] / [1 + \tan \alpha_1 \tan \alpha(AB)].$$

Assuming the angle between DA and AH is β_i , we have:

$$(104) \quad \cos \beta_i = |AH| / |AD| = |AC| \cos \alpha_i = \sin \beta \cos \alpha_i,$$

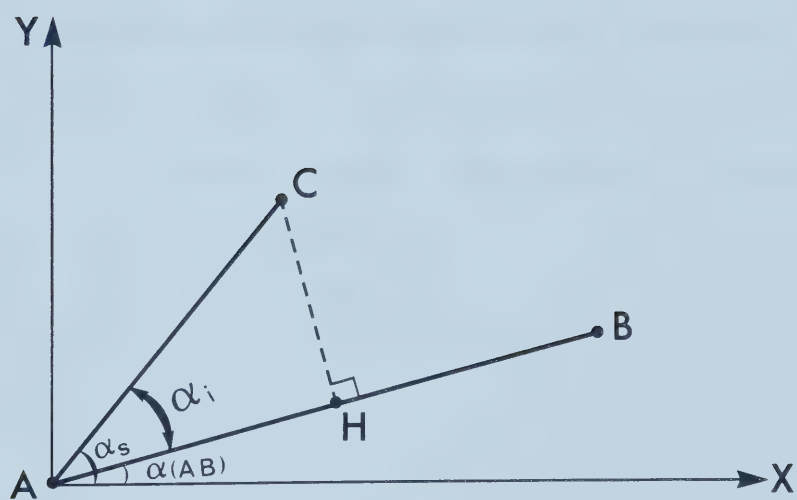
$$i = 2 \text{ or } 3,$$

$$|BB'| = \ell_i \cos \beta_i = \ell_i \sin \beta \cos \alpha_i.$$

From Fig.6 it is obvious that :

$$(105) \quad Z_i(S) = BF = |BB'| \cos \beta = 2^{-1} \ell_i \sin 2\beta \cos \alpha_i.$$

FIGURE 7. Definitions of the angles $\alpha(AB)$ and α_i and determination of the coordinate z_i .



As shown in Fig.8, because B'K is perpendicular to the x-y plane we have:

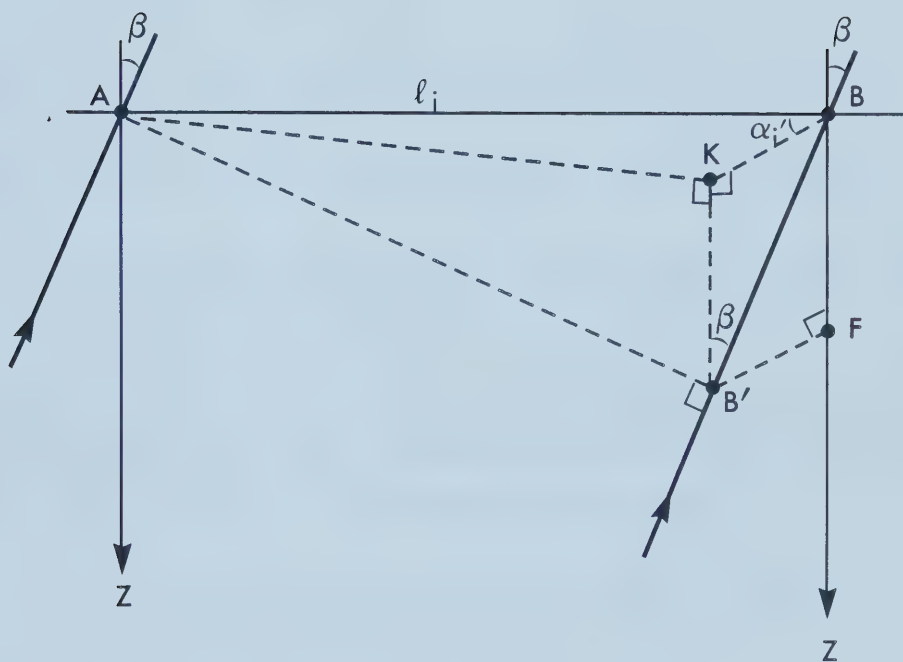
$$|BK| = |BB'| \sin \beta = \ell_i \sin^2 \beta \cos \alpha_i, \quad (106)$$

$$|AK|^2 = |AB'|^2 - |B'K|^2 = \ell_i^2 - |BB'|^2 - Z_i^2.$$

Letting the angle between BK and AB be α_i' we have:

$$\begin{aligned} (107) \quad \cos \alpha_i' &= (|BK|^2 + \ell_i^2 - |AK|^2) / (2\ell_i |BK|) \\ &= [\cos^2 \alpha_i \sin^4 \beta + \cos^2 \alpha_i \sin^2 \beta + Z_i^2(S)/\ell_i^2] \\ &\quad / [2\cos \alpha_i \sin^2 \beta]. \end{aligned}$$

FIGURE 8. Definition of the angle α_1' .



As shown in Fig.9, we have:

$$(108) \quad \cos \gamma_i = \cos[\alpha(AB) + \alpha_i'].$$

So finally, we can obtain the other coordinates of the point B' as follows:

$$(109) \quad X_i(S) = x_i - |BK| \cos \gamma_i = x_i - \ell_i \cos \alpha_i \sin^2 \beta \cos \gamma_i,$$

$$(110) \quad Y_i(S) = y_i - |BK| \sin \gamma_i = y_i - \ell_i \cos \alpha_i \sin^2 \beta \sin \gamma_i.$$

From the above geometrical relationship we have obtain the coordinats at the point B': X_i, Y_i, Z_i ($i = 2$ or 3). Also we may get $\Delta\beta$ and $\Delta\alpha$ based on Eq.(95) and (96) using the perturbed quantities. Therefore, we are now able to obtain the Gaussian coefficients as follows:

$$E(S) = [(X_2(S)-x_1)/\Delta\beta]^2 + [(Y_2(S)-y_1)/\Delta\beta]^2 + [(Z_2(S)-0)/\Delta\beta]^2,$$

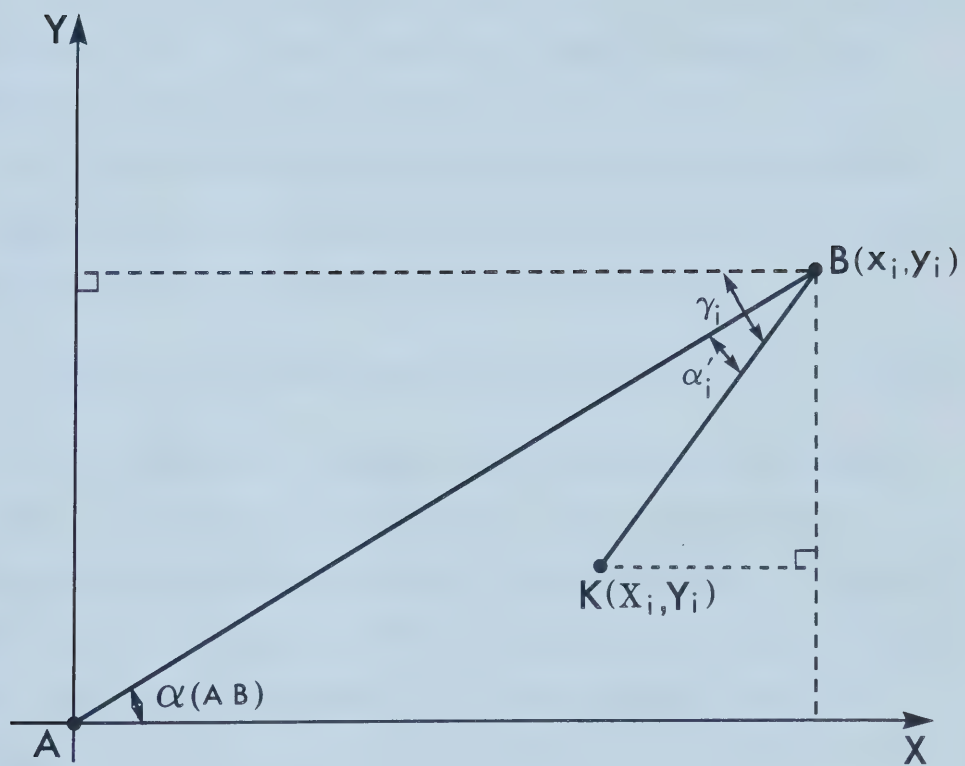
$$(111) \quad G(S) = [(X_3(S)-x_1)/\Delta\alpha]^2 + [(Y_3(S)-y_1)/\Delta\alpha]^2 + [(Z_3(S)-0)/\Delta\alpha]^2,$$

$$F(S) = [(X_2(S)-x_1)/\Delta\beta][(X_3(S)-x_1)/\Delta\alpha] + [(Y_2(S)-y_1)/\Delta\beta][(Y_3(S)-y_1)/\Delta\alpha] + [Z_2(S)/\Delta\beta][Z_3(S)/\Delta\alpha],$$

and the Jacobian, J, at the receiver point A:

$$(112) \quad J(S) = [E(S)G(S) - F(S)^2]^{1/2}.$$

FIGURE 9. Definition of the angle γ_i and determination of the coordinates X_i and Y_i .



At the taking off point (source), the coordinate, $X_i(s_0)$, $Y_i(s_0)$, $Z_i(s_0)$, where the perturbed ray intersects with the reference wavefront, could be obtained directly by solving Eq.(99) - (100). So we are able to obtain the Jacobian at the taking off point, $J(s_0)$, too.

Putting $J(S)$ and $J(s_0)$ into Eq.(42) we may determine the amplitude of the reflected ray at the receiver A when we know some of the following relative quantities:

$$\Phi(S) = R_1 R_2 \{J(s_0)\rho(s_0)V(s_0)/[J(S)\rho(S)V(S)]\}^{1/2}\Phi(s_0).$$

In the case of a medium with interfaces, we have to consider the boundary conditions at the interfaces concerning the ray path and the ray amplitude. V. Cerveny (1979) and I. Psencik (1972) have discussed them in detail. Their results are suitable for incorporation into our formulae.

10. Conclusion

From the above discussions we should obtain the following results:

(1) We may use ray path equation and the energy equation instead of the vectorial wave equation to study the kinematic and dynamic characteristics of the wave field in any weakly inhomogeneous medium, where the following restrictions are satisfied:

$$|L \nabla \rho / \rho| \ll 1,$$

$$|L \nabla \lambda / \lambda + 2\mu| \ll 1,$$

$$|L \nabla \mu / \lambda + 2\mu| \ll 1.$$

(2) The computer program ZH1 listed in Appendix I may be used to solve two point seismic ray tracing problem in a 2-D. medium and produce all kinematic characteristics of the wave field considered. The computation accuracy and efficiency of program ZH1 are highly reasonable.

(3) Based on the solution of two point ray tracing problem we described a method to compute the ray amplitudes in order to study dynamic characteristics of wave field. This would be an interesting project to follow up.

REFERENCES

Aki, K. and W. H. K. Lee (1976).

Determination of three-dimensional velocity anomalies under a seismic array using first P-arrival times from local earthquakes, Part I - A homogeneous initial model, J. Geophys. Res. 81, 4391-4399.

Aki, K., A. Christofferson, and E. S. Husebye (1977).

Determination of the three-dimensional seismic structure of the lithosphere, J. Geophys. Res. 82, 277-296.

Bjorck, A. and V. Pereyra (1970).

Solution of Vandermode systems of equations, Mathematics of computation, V.24, No.112, 893-903.

Cerveny, V., I. A. Molotkov and I. Psencik (1977).

Ray method in seismology, Charles University Press, Prague.

Cerveny, V. (1979).

Ray amplitudes of seismic body waves in laterally inhomogeneous media, Geophys. J. R. Astr. Soc., Vol. 57, 91-106.

Engdahl, E. R. and W. H. K. Lee (1976).

Relocation of local earthquakes by seismic ray tracing, J. Geophys. Res. 81, 4400-4406.

Hron, F. (1984),

Introduction to A. R. T. in seismology, Institute of Earth and Planetary Physics and Department of Physics, University of Alberta, Edmonton, Alberta.

Julian, B. R. and D. Gubbins (1977).

Three dimensional seismic ray tracing, J. Geophys. 43, 95-114.

Keller, H. B. (1969).

Accurate difference methods for linear ordinary differential systems subject to linear constraints, SIAM J. Numer. Anal., V.6, pp.8-30. MR 40 #6776.

Keller, H. B. (1974).

Accurate difference methods for nonlinear two point boundary value problems, SIAM J. Numer. Anal. 11, 305-320.

Keller, H. B. (1968).

Numerical methods for two point boundary value problems, Blaisdell, London, 184 pp.

Lentini, M. and V. Pereyra (1977).

An adaptive finite difference solver for nonlinear two point boundary problems with mild boundary layers, SIAM J. Numer. Anal. 14, 91-111.

Lentini, M. and V. Pereyra (1975).

An adaptive finite difference solver for non-linear two point boundary problems with mild boundary layers, Comp. Sc. Dept. Stanford University Rept. STAN-CS-75-530, 40 pp.

Lentini, M. and V. Pereyra (1974).

A variable order finite difference method for nonlinear multipoint boundary value problems, Math. Comp. 28, 981-1004.

Officer, C. B. (1958).

Introduction to the theory of sound transmission, McGraw Hill, New York, p.60.

Officer, C. B. (1974).

Introduction to theoretical geophysics, Chapter 6, Springer-Verlag.

Pereyra, V., W. H. K. Lee and H. B. Keller (1980).

Solving two-point seismic-ray tracing problems in a heterogeneous medium, Bull. Seism. Soc. Am. 70, 79-99.

Pereyra, V. (1973).

High order finite difference solution of differential equations, Comp. Sci. Dept. Stanford University Rept. STAN-CS-73-348.

Pereyra, V. (1968).

Iterated deferred corrections for nonlinear boundary value problems, Numer. Math. 11, 111-125.

Pereyra, V. (1967).

Iterated deferred corrections for nonlinear operator equations, Numer. Math. 10, 316-323.

Pereyra, V. and C. Sewell (1975).

Mesh selection for discrete solution of boundary problems in ordinary differential equations, Numer. Math. 23, 261-268.

Psencik, I. (1972).

Kinematics of refracted and reflected wave in inhomogeneous media with nonplanar interface, Studia Geophys. et. Geod., Vol.16, 126-152.

Tritton, D. J. (1977).

Physical fluid dynamics, New York, Reinhold.

Wesson, R. L. (1971).

Travel-time inversion for laterally inhomogeneous crustal velocity models, Bull. Seism. Soc. Am. 61, 729-746.

Yang, J. P. and W. H. K. Lee (1976).

Preliminary investigations on computational methods for

solving two point seismic ray tracing problems in a
heterogeneous and isotropic medium, U. S. Geol. Surv.,
Open-File Rept. 76-707, 66 pp.

APPENDIX I


```

1  /COMPILE SIZE=60P
2  BEGIN
3  INTEGER M, N;
4  STRING DATA;
5  REAL Z0, Z1, Z2, ALF1, ALF2, ALF3, ALF4, TOL;
6  READ(Z0, Z1, Z2, ALF1, ALF2, ALF3, ALF4);
7  M := 2;
8  N := 21;
9  TOL:= 1.0' -3;
10 BEGIN
11 REAL A, B, EPSMAC, EPS, ERROLD, ERRNEW, C1, HMAX,
12     RABS, REOLD, SUM, TEM, AUX1, U1;
13 INTEGER KJH, JERROR, I, J, NUK, KMAX, N1, I1, IERROR, INWT,
14     K1, NU2, K, NUINT, NO2, NO21, L, JK, AEMEK;
15 REAL ARRAY ABT, ALPHA(1::M);
16 REAL ARRAY A1, B1(1::M, 1::M);
17 REAL ARRAY AA, BB, CC(1::50);
18 LOGICAL DIVNEW;
19 PROCEDURE FF(REAL ARRAY X(*); REAL ARRAY Y(*,*);
20             INTEGER VALUE N; REAL ARRAY F(*,*));
21 BEGIN
22 FOR I := 1 UNTIL N DO
23 BEGIN
24 F(1,I) := Y(2,I);
25 F(2,I) := (1+Y(2,I)**2) * (Z1*Y(2,I)-Z2)
26           /(Z0+Z1*X(I)+Z2*Y(1,I));
27 END;
28 END FF;
29 PROCEDURE COEGE(INTEGER VALUE N, NP; REAL ARRAY C, BB(*);
30               REAL VALUE E);
31 BEGIN
32 REAL ARRAY ALF(1::50);
33 INTEGER I, NN, LL, J, K, KM1, JM1;
34 REAL XKIN;
35 FOR I := 1 UNTIL N DO
36 C(I) := BB(I);
37 FOR I := 1 UNTIL N DO
38 ALF(I) := I - NP - E;
39 NN := N - 1;
40 FOR I := 1 UNTIL NN DO
41 BEGIN
42 LL := N - I;
43 FOR J := 1 UNTIL LL DO
44 BEGIN
45 K := N - J + 1;
46 C(K) := C(K) - ALF(I) * C(K-1)
47 END;
48 END;
49 FOR I := 1 UNTIL NN DO
50 BEGIN
51 K := N - I;
52 XKIN := 1.0/K;
53 KM1 := K + 1;
54 FOR J := KM1 UNTIL N DO
55 BEGIN
56 C(J) := C(J) * XKIN;
57 JM1 := J - 1;
58 C(JM1) := C(JM1) - C(J);
59 END;
60 END;

```



```

61  END COEGE;
62  PROCEDURE U2DCGS (INTEGER VALUE K, P, Q, N, M; REAL ARRAY
63      A(*); REAL ARRAY Y, S(*,*); REAL VALUE EE;
64      PROCEDURE COEGEN; INTEGER RESULT IERROR);
65  BEGIN
66  REAL ARRAY C(1::50);
67  INTEGER I, KK1, KK, KMID, KMID1, L, IT, NF, II, J, KMIDP1;
68  REAL ACUM;
69  IF (K > (N + 1 - Q) DIV P) OR (P < 1) OR (K < 1) THEN GOTO S100;
70  IF (Q = 0) THEN GOTO S10;
71  FOR I := 1 UNTIL Q DO A(I) := 0;
72  S10: KK1 := Q + P * K;
73  KK := KK1 - 1;
74  KMID := KK1 DIV 2;
75  IERROR := 0;
76  KMID1 := KMID - 1;
77  IF (KMID < 1) THEN GOTO S25;
78  FOR I := 1 UNTIL KMID1 DO
79  BEGIN
80  COEGEN(KK1, I, C, A, EE);
81  FOR L := 1 UNTIL M DO
82  BEGIN
83  ACUM := 0;
84  FOR J := 1 UNTIL KK1 DO
85  ACUM := ACUM + C(J) * Y(L, J);
86  S(L, I) := ACUM;
87  END;
88  END;
89  S25: COEGEN(KK1, KMID, C, A, EE);
90  NF := N + 1 - KK1 + KMID;
91  FOR I := KMID UNTIL NF DO
92  BEGIN
93  II := I - KMID;
94  FOR L := 1 UNTIL M DO
95  BEGIN
96  ACUM := 0;
97  FOR J := 1 UNTIL KK1 DO
98  ACUM := ACUM + C(J) * Y(L, II + J);
99  S(L, I) := ACUM;
100  END;
101  END;
102  KMIDP1 := KMID + 1;
103  FOR I := KMIDP1 UNTIL KK DO
104  BEGIN
105  COEGEN(KK1, I, C, A, EE);
106  II := N - KK;
107  FOR L := 1 UNTIL M DO
108  BEGIN
109  ACUM := 0;
110  FOR J := 1 UNTIL KK1 DO
111  ACUM := ACUM + C(J) * Y(L, II + J);
112  S(L, I + II) := ACUM;
113  END;
114  END;
115  GOTO S20;
116  S100 : IERROR := 1;
117  S20:
118  END U2DCGS;
119  PROCEDURE JACOB (REAL VALUE XX; REAL ARRAY YY(*);
120      REAL ARRAY XJAC(*,*));

```



```

121 BEGIN
122 XJAC(1,1) := 0;
123 XJAC(1,2) := 1.0;
124 XJAC(2,1) := (1+YY(2)**2)*(Z2**2 - Z1*Z2*YY(2))
125             /(Z0+Z1*XX + Z2*YY(1))**2;
126 XJAC(2,2) := (Z1-2*Z2*YY(2)+3*Z1*YY(2)**2)
127             /(Z0+Z1*XX+Z2*YY(1));
128 END JACOB;
129 PROCEDURE GS( INTEGER VALUE M,N; REAL ARRAY A(*,*) );
130 BEGIN
131   INTEGER IO, JO, NM, I, J, K, TI;
132   REAL C, T;
133   INTEGER ARRAY Z(1::M);
134   NM := N + M;
135   FOR K := 1 UNTIL M DO Z(K) := K;
136   FOR K := 1 UNTIL M DO
137     BEGIN
138       C := 0;
139       FOR I := K UNTIL M DO
140         FOR J := K UNTIL M DO
141           IF ( ABS(A(I,J)) > ABS(C) ) THEN
142             BEGIN
143               C := A(I,J);
144               IO := I;
145               JO := J
146             END;
147           IF ( ABS(C) < 1.0E-20 ) THEN
148             BEGIN
149               WRITE ( "SMOLLER" );
150               GOTO SEND;
151             END;
152           IF ( IO = K ) THEN GOTO L1;
153           FOR J := 1 UNTIL NM DO
154             BEGIN
155               T := A(K,J);
156               A(K,J) := A(IO,J);
157               A(IO,J) := T
158             END;
159           L1: IF (JO = K ) THEN GOTO L2;
160           FOR I := K UNTIL M DO
161             BEGIN
162               T := A(I,K);
163               A(I,K) := A(I,JO);
164               A(I,JO) := T
165             END;
166           TI := Z(K);
167           Z(K) := Z(JO);
168           Z(JO) := TI;
169           L2: C := 1/C;
170           FOR J := K+1 UNTIL NM DO
171             A(K,J) := A(K,J) * C;
172           FOR I := K+1 UNTIL M DO
173             FOR J := K+1 UNTIL NM DO
174               A(I,J) := A(I,J) - A(I,K) * A(K,J)
175             END;
176           FOR I := M + 1 UNTIL NM DO
177             FOR K := M - 1 STEP -1 UNTIL 1 DO
178               BEGIN
179                 C := A(K,I);
180                 FOR J := K + 1 UNTIL M DO

```



```

181 C := C - A(K,J) * A(J,I);
182 A(K,I) := C
183 END;
184 FOR K:= 1 UNTIL M DO
185 BEGIN
186 FOR I := K + 1 UNTIL M DO
187 IF ( Z(I) = K ) THEN
188 BEGIN
189 FOR J := M + 1 UNTIL NM DO
190 BEGIN
191 T := A(K,J);
192 A(K,J) := A(I,J);
193 A(I,J) := T;
194 END;
195 Z(I) := Z(K);
196 GOTO L3
197 END;
198 L3:
199 END
200 END GS;
201 PROCEDURE SYSLIN( INTEGER VALUE M,N; REAL ARRAY X(*);
202 REAL ARRAY Y(*,*); REAL ARRAY H(*); REAL ARRAY
203 RES, A1, B1(*,*) ; PROCEDURE JACOB, GS2; REAL
204 ARRAY UU(*,*) );
205 BEGIN
206 INTEGER I, J, M2, MPN, M1, K1, K, L;
207 REAL X1, XL, SUM;
208 REAL ARRAY H2(1::N-1);
209 REAL ARRAY Y1, U(1::M);
210 REAL ARRAY JAB, S, R(1::M, 1::M);
211 REAL ARRAY T, VM, AUX, AB1(1::M,1::M+1);
212 REAL ARRAY A(1::M,1::2 * M + 1);
213 REAL ARRAY V(1::N * M, 1::M + 1);
214 REAL ARRAY U1(1::M * N);
215 X1 := X(1);
216 FOR I := 1 UNTIL M DO
217 Y1(I) := Y(I,1);
218 JACOB(X1, Y1, JAB);
219 M2 := M * M;
220 MPN := M * N;
221 M1 := M + 1;
222 FOR I := 1 UNTIL M DO
223 T(I,M1) := RES(I,2);
224 FOR L := 2 UNTIL N DO
225 BEGIN
226 FOR J := 1 UNTIL M DO
227 U(J) := Y(J, L);
228 H2(L - 1) := 0.5 * H(L - 1);
229 FOR J := 1 UNTIL M DO
230 FOR I := 1 UNTIL M DO
231 BEGIN
232 S(I,J) := H2(L - 1) * JAB(I,J);
233 IF (I=J) THEN S(I,J) := S(I,J) + 1.0;
234 END;
235 XL := X(L);
236 JACOB(XL, U, JAB);
237 FOR I := 1 UNTIL M DO
238 FOR J := 1 UNTIL M DO
239 BEGIN
240 R(I,J) := -H2(L-1) * JAB(I,J);

```



```

241 IF (I=J) THEN R(I,J) := R(I,J) + 1.0;
242 END;
243 IF (L > 2) OR (L < 2) THEN GOTO S1300;
244 FOR I := 1 UNTIL M DO
245   FOR J := 1 UNTIL M DO
246     T(I,J) := - S(I,J);
247   GOTO S1700;
248 S1300 :
249   FOR K1 := 1 UNTIL M1 DO
250     FOR I := 1 UNTIL M DO
251       BEGIN
252         SUM := 0;
253         FOR J := 1 UNTIL M DO
254           SUM := SUM + S(I,J) * VM(J,K1);
255         T(I,K1) := SUM;
256       END;
257     FOR I := 1 UNTIL M DO
258       T(I,M1) := T(I,M1) + RES(I,L);
259   S1700:
260   FOR I := 1 UNTIL M DO
261     FOR J := 1 UNTIL M DO
262       A(I,J) := R(I,J);
263     FOR J := 1 UNTIL M1 DO
264       FOR I := 1 UNTIL M DO
265         A(I,M + J) := T(I,J);
266       GS2(M, M1, A);
267     FOR J := 1 UNTIL M1 DO
268       FOR I := 1 UNTIL M DO
269         VM(I,J) := A(I,M+J);
270     FOR J := 1 UNTIL M1 DO
271       FOR I := 1 UNTIL M DO
272         V((L-1) * M + I, J) := VM(I,J);
273     END;
274     FOR J := 1 UNTIL M1 DO
275       FOR I := 1 UNTIL M DO
276         BEGIN
277           SUM := 0;
278           FOR K := 1 UNTIL M DO
279             SUM := SUM + B1(I,K) * VM(K,J);
280           AUX(I,J) := SUM;
281         END;
282     FOR J := 1 UNTIL M DO
283       FOR I := 1 UNTIL M DO
284         R(I,J) := A1(I,J) - AUX(I,J);
285     FOR I := 1 UNTIL M DO
286       U1(I) := RES(I,1) - AUX(I,M1);
287     FOR I := 1 UNTIL M DO
288       FOR J := 1 UNTIL M DO
289         AB1(I,J) := R(I,J);
290     FOR I := 1 UNTIL M DO
291       AB1(I,M + 1) := U1(I);
292     GS2(M, 1, AB1);
293     FOR I := 1 UNTIL M DO
294       U1(I) := AB1(I, M + 1);
295     FOR I := M1 UNTIL MPN DO
296       BEGIN
297         SUM := V(I,M1);
298         FOR J := 1 UNTIL M DO
299           SUM := SUM - V(I,J) * U1(J);
300       U1(I) := SUM;

```



```

301 END;
302 FOR J := 1 UNTIL N DO
303 BEGIN
304 K1 := (J - 1) * M;
305 FOR I := 1 UNTIL M DO
306 UU(I,J) := U1(K1 + I);
307 END
308 END;
309 NEWLINE(13);
310 WRITE("                                VELOCITY MODEL:");
311 NEWLINE(2);
312 WRITE("                                V=Z0+Z1*X+Z2*Z");
313 IOCONTROL("SPACE");
314 WRITE("                                Z0=",Z0);
315 WRITEON("Z1=",Z1);
316 WRITEON("Z2=",Z2);
317 NEWLINE(2);
318 WRITE("                                X(0)=",ALF1);
319 WRITEON("Z(0)=",ALF2);
320 WRITE("                                X(1)=",ALF3);
321 WRITEON("Z(1)=",ALF4);
322 IOCONTROL("EJECT");
323 A := ALF1;
324 B := ALF3;
325 EPSMAC := 1.0' -40;
326 JERROR := 0;
327 DIVNEW := FALSE;
328 KJH := 0;
329 NUK := 0;
330 FOR I := 1 UNTIL M DO
331 BEGIN
332 ALPHA(I) := 0;
333 FOR J := 1 UNTIL M DO
334 BEGIN
335 A1(I,J) := 0;
336 B1(I,J) := 0;
337 END;
338 END;
339 A1(1,1) := 1.0;
340 B1(2,1) := 1.0;
341 ALPHA(1) := ALF2;
342 ALPHA(2) := ALF4;
343 AEMEK := 0;
344 JK := -10;
345 LLL3:
346 JK := JK + 10;
347 BEGIN
348 REAL ARRAY H(1::N-1);
349 REAL ARRAY X(1::N);
350 REAL ARRAY Y, RES(1::M, 1::N);
351 INTEGER ARRAY IQJ(1::N-1);
352 PROCEDURE ENDPU;
353 BEGIN
354 REAL ARRAY SSUM,TSUM(1::N);
355 SSUM(1) := 0.0;
356 TSUM(1) := 0.0;
357 FOR I := 2 UNTIL N DO
358 BEGIN
359 U1 := SQRT((X(I)-X(I-1))**2+(Y(1,I)-Y(1,I-1))**2);
360 SSUM(I) := SSUM(I-1)+U1;

```



```

361      TSUM(I) := TSUM(I-1)+U1/(Z0+0.5*Z1*(X(I)+X(I-1))+
362          0.5*Z2*(Y(1,I)+Y(1,I-1)));
363      END;
364      NEWLINE(3);
365      R_W := 10;
366      I_W := 3;
367      WRITE("                N=", N);
368      WRITE("                TOL=", TOL, "(", ERRNEW, ")");
369      NEWLINE(1);
370      S_W := 5;
371      R_D := 3;
372      R_FORMAT := "F";
373      K1 := (N-1) DIV 20;
374      FOR J := 1 STEP K1 UNTIL N DO
375      BEGIN
376          WRITE("                N=", J);
377          WRITEON("X=", X(J));
378          WRITEON("Z=", Y(1,J));
379          SUM := ARCTAN(Y(2,J))*180/3.14159265;
380          WRITEON("Q=", SUM);
381          WRITEON("T=", TSUM(J));
382          WRITEON("S=", SSUM(J));
383      END;
384      K1 := (N-1) REM K1;
385      IF (K1=0) THEN GOTO SENDPU;
386      WRITE("                N=", N);
387      WRITEON("X=", X(N));
388      WRITEON("Z=", Y(1,N));
389      SUM := ARCTAN(Y(2,N))*180/3.14159265;
390      WRITEON("Q=", SUM);
391      WRITEON("T=", TSUM(N));
392      WRITEON("S=", SSUM(N));
393      SENDPU:
394      END;
395      K := 0;
396      BEGIN
397      REAL ARRAY F, UU(1::M, 1::N);
398      REAL ARRAY SK(1::M, 1::N-1);
399      REAL HH;
400      IF DIVNEW THEN GOTO S40;
401      IF (KJH >= 1) THEN GOTO S120;
402      S40:
403      TEM := (ALF2 - ALF4)/(ALF3 - ALF1);
404      SUM := (ALF3 - ALF1)/(N-1);
405      FOR J := 1 UNTIL N DO
406      BEGIN
407          Y(1,J) := ALF2 - TEM * SUM*(J-1);
408          Y(2,J) := -TEM;
409      END;
410      BB(1) := 1.0;
411      FOR I := 2 UNTIL 50 DO BB(I) := 0;
412      HH := (B-A)/(N-1);
413      IF (EPSMAC > 0.01 * HH ** 4) THEN EPS := EPSMAC
414      ELSE EPS := 0.01 * HH ** 4;
415      NUK := 0;
416      X(1) := A;
417      X(N) := B;
418      FOR I := 2 UNTIL N-1 DO
419      X(I) := A + (I-1) * HH;
420      FOR I := 1 UNTIL N-1 DO

```



```

421 H(I) := HH;
422 S120:
423 ERROLD := 1.0;
424 C1 := 0.8;
425 FOR I := 1 UNTIL M DO
426 FOR J := 1 UNTIL N-1 DO
427 SK(I,J) := 0;
428 IF (KJH >= 1) THEN
429 BEGIN
430 REWIND(7);
431 FOR J := 1 UNTIL N DO
432 FOR I := 1 UNTIL M DO
433 GET(7, NULL, Y(I,J));
434 FOR I := 1 UNTIL N-1 DO
435 GET(7, NULL, H(I));
436 RELEASE(7);
437 X(1) := A;
438 X(N) := B;
439 FOR I := 2 UNTIL N-1 DO
440 X(I) := X(I-1) + H(I-1);
441 END;
442 KMAX := (N-2) DIV 2;
443 N1 := N-1;
444 HMAX := 1/N1;
445 IF (NUK = 0) THEN GOTO S405;
446 FF(X, Y, N, F);
447 U2DCGS(NUK, 2, 2, N1, M, AA, F, RES, 0.5, COEGE, IERROR);
448 FOR J := 1 UNTIL N1 DO
449 FOR I := 1 UNTIL M DO
450 SK(I,J) := H(J) * RES(I,J);
451 S405:
452 INWT := 0;
453 REOLD := 1.0;
454 DIVNEW := FALSE;
455 IF (KMAX > NUK + 1) THEN GOTO S410;
456 NUK := NUK + 1;
457 GOTO SPAS;
458 S410:
459 RABS := 0;
460 FOR I := 1 UNTIL M DO
461 BEGIN
462 SUM := ALPHA(I);
463 FOR J := 1 UNTIL M DO
464 SUM := SUM - A1(I,J) * Y(J,1) - B1(I,J) * Y(J,N);
465 RES(I,1) := SUM;
466 TEM := ABS(RES(I,1));
467 IF (TEM > RABS) THEN RABS := TEM;
468 END;
469 FF(X, Y, N, F);
470 FOR I := 2 UNTIL N DO
471 FOR J := 1 UNTIL M DO
472 BEGIN
473 RES(J,I) := -Y(J,I) + Y(J,I-1) + H(I-1)/2 * (F(J,I)
474 + F(J,I-1)) + SK(J,I-1);
475 TEM := ABS(RES(J,I));
476 IF (TEM > RABS) THEN RABS := TEM;
477 END;
478 IF (RABS > 1.0) AND (INWT=1) THEN
479 BEGIN
480 DIVNEW := TRUE;

```



```

481 N := N+10;
482 GOTO LLL3;
483 END;
484 IF (RABS>REOLD) AND (INWT>1) THEN GOTO S1500;
485 IF (INWT>=3) THEN GOTO S1500;
486 SYSLIN(M, N, X, Y, H, RES, A1, B1, JACOB, GS, UU);
487 REOLD := RABS;
488 FOR I := 1 UNTIL M DO
489   FOR J := 1 UNTIL N DO
490     Y(I,J) := Y(I,J) + UU(I,J);
491   INWT := INWT + 1;
492 GOTO S410;
493 S1500:
494 NUK := NUK + 1;
495 NU2 := 2 * NUK + 1;
496 AA(NU2) := -NUK / (2 ** (2 * NUK - 1) * NU2);
497 AA(NU2 + 1) := 0;
498 U2DCGS(NUK, 2, 2, N1, M, AA, F, RES, 0.5, COE, IERR);
499 IF (IERR = 1) THEN GOTO SPAS;
500 FOR J := 1 UNTIL N1 DO
501   FOR I := 1 UNTIL M DO
502     BEGIN
503       AUXI := RES(I,J) * H(J);
504       RES(I,J) := SK(I,J) - AUXI;
505       SK(I,J) := AUXI;
506     END;
507 SYSLIN(M, N, X, Y, H, RES, A1, B1, JACOB, GS, UU);
508 ERRNEW := 0;
509 FOR I := 1 UNTIL M DO
510   ABT(I) := 0.0;
511   FOR J := 1 UNTIL N DO
512     FOR I := 1 UNTIL M DO
513       BEGIN
514         U1 := ABS(UU(I,J));
515         IF (U1 > ABT(I)) THEN ABT(I) := U1;
516       END;
517   FOR J := 1 UNTIL M DO
518     BEGIN
519       IF (ABT(J) > ERRNEW) THEN ERRNEW := ABT(J);
520     END;
521   IF (ERRNEW <= TOL) THEN
522     BEGIN
523       ENDP;
524       GOTO SEND;
525     END;
526   CC(NUK) := ERRNEW;
527   IF (ERRNEW <= 0.1 * EROLD) THEN GOTO S2550;
528   IF (ERRNEW > C1 * EROLD) THEN GOTO SPAS;
529   C1 := 0.5 * C1;
530 S2550:
531 EROLD := ERRNEW;
532 IF (EPS = EPSMAC) THEN
533   BEGIN
534     WRITE("EROLD=", EROLD);
535     WRITE("NEED CHANGE TOL");
536     STOP(NULL);
537   END;
538 IF (EPSMAC > 0.01 * HMAX ** 2 * EROLD) THEN
539   EPS := EPSMAC ELSE EPS := 0.01 * HMAX ** 2 * EROLD;
540 GOTO S405;

```



```

541 SPAS:
542 IF (JK = 140) THEN
543 BEGIN
544 WRITE("ERROLD=", ERROLD);
545 WRITE("NEED CHANGE TOL");
546 GOTO SEND;
547 END;
548 K := NUK - 1;
549 I1 := 50;
550 GOTO AEME;
551 END;
552 AEME:
553 BEGIN
554 REAL ARRAY EJ(1::N-1);
555 INTEGER IQ,NMAX,IQSR;
556 REAL SAKM, BMA, ELEV,LMTA,ESR,ALG;
557 SIQ:
558 TEM := TOL;
559 SAKM := 1/(2*K+2);
560 ELEV := TEM ** ROUND(SAKM);
561 U1 := 0;
562 FOR J := 1 UNTIL N1 DO
563 BEGIN
564 AUXI := H(J) ** (2*K+2);
565 SUM := 0;
566 FOR I := 1 UNTIL M DO
567 BEGIN
568 TEM := ABS(RES(I,J));
569 IF (TEM > SUM) THEN SUM := TEM;
570 END;
571 SUM := SUM/AUXI;
572 IF (SUM > U1) THEN U1 :=SUM;
573 END;
574 LMTA := U1/I1 ** (2*K+2);
575 FOR J := 1 UNTIL N1 DO
576 BEGIN
577 AUXI := H(J) ** (2*K+2);
578 SUM := 0;
579 FOR I := 1 UNTIL M DO
580 BEGIN
581 TEM := ABS(RES(I,J));
582 IF (TEM > SUM) THEN SUM := TEM;
583 END;
584 SUM := SUM / AUXI;
585 IF (SUM >= LMTA) THEN U1 := SUM ** ROUND(SAKM);
586 IF (LMTA > SUM) THEN U1 := LMTA ** ROUND(SAKM);
587 EJ(J) := H(J) * U1 ;
588 END;
589 FOR J := 1 UNTIL N1 DO
590 IQJ(J) := ROUND(EJ(J) / ELEV);
591 IQ := 0;
592 FOR J := 1 UNTIL N1 DO
593 BEGIN
594 IF (IQJ(J) = 0) THEN IQ :=IQ + IQJ(J) ELSE
595 IQ := IQ + IQJ(J) - 1;
596 END;
597 NMAX := 2000 DIV M;
598 U1 := NMAX;
599 ALG := 1.4;
600 RABS := NMAX - N;

```



```

601 IF (RABS > 70.0) THEN RABS := 70.0;
602   SJU:
603   IF (IQ < 0.04 * N) THEN GOTO SLS;
604   IF (IQ > RABS) THEN GOTO SGR;
605   IF (IQ >= 0.04 * N) AND (IQ <= RABS) THEN
606   BEGIN
607     AEMEK := K;
608     GOTO S2600;
609   END;
610   SLS:
611   SUM := 1.4 * N;
612   IF (SUM < U1) THEN TEM := SUM ELSE TEM := U1;
613   SUM := TEM - N - 1;
614   IQSR := ROUND(SUM);
615   ESR := (N+IQ+1) / (N+IQSR+1) * ELEV;
616   FOR J := 1 UNTIL N1 DO
617     IQJ(J) := ROUND(EJ(J) / ESR);
618   IQ := 0;
619   FOR J := 1 UNTIL N1 DO
620   BEGIN
621     IF (IQJ(J) = 0) THEN IQ := IQ+IQJ(J) ELSE
622     IQ := IQ + IQJ(J) - 1;
623   END;
624   IF (IQ < 0.04 * N) THEN
625   BEGIN
626     IF (K > 0) THEN K := K-1;
627     IF (K = 0) THEN
628     BEGIN
629       IF ((2*N - 1) <= NMAX) THEN GOTO S026 ELSE
630       BEGIN
631         WRITE("NMAX SO SMALL THAT STOP");
632         GOTO SEND;
633       END;
634     END;
635     IF (K > 0) THEN GOTO SIQ;
636   END;
637   IF (IQ >= 0.04 * N) THEN GOTO SJU;
638   SGR:
639   SUM := ALG * N;
640   IF (SUM < U1) THEN TEM := SUM ELSE
641   TEM := U1;
642   SUM := TEM - N - 1;
643   IQSR := ROUND(SUM);
644   ESR := (N+IQ+1) / (N+IQSR+1) * ELEV;
645   FOR J := 1 UNTIL N1 DO
646     IQJ(J) := ROUND((EJ(J) / ESR));
647   IQ := 0;
648   FOR J := 1 UNTIL N1 DO
649   BEGIN
650     IF ( IQJ(J) = 0) THEN IQ := IQ + IQJ(J) ELSE
651     IQ := IQ + IQJ(J) -1;
652   END;
653   IF (IQ > RABS) THEN
654   BEGIN
655     ALG := ALG - 0.1;
656     IF (ALG >= 1.1) THEN GOTO SGR;
657   END;
658   IF (IQ <= RABS) THEN GOTO SJU;
659   END;
660   S2600:

```



```

661 BEGIN
662 REAL ARRAY SK(1::M,1::N-1);
663 I1 := 1;
664 FOR I := 1 UNTIL N1 DO
665 BEGIN
666 IF (IQJ(I) > 0) THEN
667 I1 := I1 + 1 + IQJ(I) - 1 ELSE
668 I1 := I1 + 1;
669 END;
670 BEGIN
671 REAL ARRAY Y1(1::M,1::I1);
672 REAL ARRAY H1(1::I1-1);
673 REAL ARRAY CC1(1::50);
674 REAL HH;
675 INTEGER J1;
676 KJH := KJH + 1;
677 NUINT := NUK;
678 FOR I := 1 UNTIL M DO
679 FOR J := 1 UNTIL N DO
680 RES(I,J) := Y(I,J);
681 N1 := N - 1;
682 L := 1;
683 FOR J := 1 UNTIL N1 DO
684 BEGIN
685 IF (IQJ(J) > 1) THEN
686 BEGIN
687 K1 := L;
688 L := L + IQJ(J) - 1;
689 FOR I := 1 UNTIL M DO Y1(I,K1) := RES(I,J);
690 FOR J1 := K1+1 UNTIL L DO
691 BEGIN
692 U1 := 1/IQJ(J) * (J1-K1);
693 U2DCGS(NUINT,2,0,N1,M,BB,Y,SK,U1,COEGE,IERROR);
694 FOR I := 1 UNTIL M DO
695 Y1(I,J1) := SK(I,J);
696 END;
697 END;
698 IF (IQJ(J) <= 1) THEN
699 BEGIN
700 FOR I := 1 UNTIL M DO
701 Y1(I,L) := RES(I,J);
702 END;
703 L := L+1;
704 END;
705 FOR I := 1 UNTIL M DO
706 Y1(I,I1) := RES(I,N);
707 L := 1;
708 FOR I := 1 UNTIL N1 DO
709 BEGIN
710 IF (IQJ(I) > 1) THEN
711 BEGIN
712 HH := H(I) / IQJ(I);
713 K1 := L;
714 L := L+IQJ(I)-1;
715 FOR J := K1 UNTIL L DO
716 H1(J) := HH;
717 END;
718 IF (IQJ(I) <= 1) THEN
719 H1(L) := H(I);
720 L := L + 1;

```



```

721 END;
722 ASSIGN(7,"DATA");
723 FOR J := 1 UNTIL I1 DO
724   FOR I := 1 UNTIL M DO
725     PUT(7, NULL, Y1(I,J));
726   FOR I := 1 UNTIL (I1 - 1) DO
727     PUT(7, NULL, H1(I));
728   PUT(7, NULL, 0.0);
729   FOR I := 1 UNTIL NUK DO
730     CC1(I) := CC(I);
731   K1 := 0;
732   FOR I := 1 UNTIL N1 DO
733     IF (IQJ(I) > K1) THEN K1 := IQJ(I);
734   NUK := 0;
735   IF (EPSMAC > 0.01 * (1/(I1-1)) ** 4) THEN
736     EPS := EPSMAC ELSE EPS := 0.01 * (1/(I1-1)) ** 4;
737   N := I1;
738   GOTO LLL3;
739 END;
740 END;
741   S026:
742   BEGIN
743     REAL ARRAY Y1(1::M,1::2*N-1);
744     REAL ARRAY SK(1::M,1::N);
745     REAL ARRAY H1(1::2*(N-1));
746     REAL ARRAY CC1(1::50);
747     KJH := KJH + 1;
748     HMAX := H(N1);
749     FOR I := 1 UNTIL N-1 DO
750       BEGIN
751         TEM := ABS(H(I));
752         IF (TEM > HMAX) THEN HMAX := TEM;
753       END;
754     NUINT := NUK;
755     FOR I := 1 UNTIL NUK DO
756       CC1(I) := CC(I);
757     K := 0;
758     FOR I := 1 UNTIL NUK DO
759       BEGIN
760         K := K-1;
761         CC1(I) := CC1(I) * 4.0 ** K;
762       END;
763     IF (ERRNEW < ERROLD) THEN ERROLD := ERRNEW;
764     K := 0;
765     FOR I := 1 UNTIL NUK DO
766       BEGIN
767         K := K+1;
768         IF (ERROLD < CC1(I)) THEN GOTO S270;
769       GOTO S275;
770     S270:
771     END;
772     S275:
773     NUK := K - 1;
774     IF (EPSMAC > 0.005*HMAX**2*ERROLD) THEN EPS := EPSMAC
775     ELSE EPS := 0.005 * HMAX ** 2 * ERROLD;
776     FOR I := 1 UNTIL M DO
777       FOR J := 1 UNTIL N DO
778         RES(I,J) := Y(I,J);
779     N := 2 * N - 1;
780     N02 := (N+1) DIV 2;

```



```

781      N021 := N02 - 1;
782      U2DCGS(NUINT,2,0,N021,M,BB,Y,SK,0.5,COEGE,IERROR);      98
783      FOR J := 1 UNTIL N021 DO
784      FOR I := 1 UNTIL M DO
785      BEGIN
786      Y1(I,2*J-1) := RES(I,J);
787      Y1(I,2*J) := SK(I,J);
788      END;
789      FOR L := 1 UNTIL M DO
790      Y1(L,N) := RES(L,N02);
791      FOR I := 1 UNTIL N021 DO
792      BEGIN
793      H1(2*I-1) := H(I)/2;
794      H1(2*I) := H(I)/2;
795      END;
796      ASSIGN(7,"DATA");
797      FOR J := 1 UNTIL N DO
798      FOR I := 1 UNTIL M DO
799      PUT(7,NULL,Y1(I,J));
800      FOR I := 1 UNTIL N-1 DO
801      PUT(7,NULL,H1(I));
802      PUT(7,NULL,0.0);
803      GOTO LLL3;
804      END;
805      END;
806      SEND:
807      END;
808      END.

```

End of file

University of Alberta Library



0 1620 0682 2207

B34328